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EARTH STRAIN MEASUREMENTS  
WITH THE TRANSPORTABLE LASER RANGING SYSTEM:  
FIELD TECHNIQUES AND PLANNING

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## ABSTRACT

We have conducted a feasibility study to examine the potential of the Transportable Laser Ranging System (TLRS) for monitoring the ground deformation around satellite ranging stations and other geodetic control points. Emphasis has been placed on testing the usefulness of the relative lateration technique. The temporal variation of the ratio of the length of each survey line to the mean length of all survey lines in a given area is directly related to the mean shear strain rate for the area. The data from a series of experimental measurements taken over the Los Angeles basin from a TLRS station at Mt. Wilson show that such ratios can be determined to an accuracy of one part in  $10^7$  with a measurement program lasting for three days and without using any corrections for variations in atmospheric conditions. A numerical experiment using a set of hypothetical data indicates that reasonable estimates of the present shear strain rate and the direction of the principal axes in southern California can be deduced from such measurements over an interval of one to two years. Thus, the relative lateration from the TLRS appears to be a very economical way to monitor ground deformations, although there has been no opportunity yet to measure the actual ground strain by reoccupying the Mt. Wilson site.

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## I. INTRODUCTION

With the recent development in ground-to-satellite laser ranging and Very Long Baseline Interferometry (VLBI) techniques, it is now possible to measure precisely distances between locations separated by several hundreds to thousands of kilometers. This makes it possible to monitor relative movements of globally distributed points on the earth for geodynamic studies. However, one question that must be answered is how representative each of these positions thus occupied is for the region in which it is located. If some of these locations are experiencing localized movements which are not representative of the region, the global measurement would give erroneous results. An answer to this questions can be found by measuring regional deformations around each location.

A conventional method for determining regional deformations is to perform repeated survey using an electro-optical distance measuring (EDM) device (e.g. Savage et al. [1981]). However, such surveys are expensive, and are rather limited in range. We, therefore, have looked for a better alternative. The development of the Transportable Laser Ranging System (TLRS) for ground-to-LAGEOS (Laser Geodynamics Earth Orbiting Satellite) ranging [Silverberg and Byrd, 1981] has given us an opportunity to test such an alternative. Because of its high sensitivity, being capable of detecting single photon returns, the TLRS can measure distances to small targets (retro-reflectors) at any visible points much beyond the normal ranges of other EDM devices. Thus, this system may provide economical measurements of strain fields in areas more than 200 km in diameter. If successful, such measurements will be valuable not only in the immediate neighborhood of satellite ranging stations, but also in understanding the dynamic behavior of both plate boundaries and areas internal to plates.

We have conducted a limited feasibility study to examine this potential. Although the TLRS is a powerful system, it also has certain limitations when used for a ground-to-ground ranging. The most important is the uncertainty of measurement results due to variability in atmospheric conditions. To bypass this problem and avoid the expense of flying an aircraft to monitor the atmospheric conditions along the path of the laser beam, we have examined the use of the relative lateralization, or the ratio method, which was used earlier by Carter and Vincenty [1978] in an experimental survey around the McDonald Observatory.

We originally planned repeated field experiments at several sites in the western United States. However, because of many scheduling conflicts and delays associated with the overall TLRS-LAGEOS ranging experiments, the only field experiment we could perform during the current contract was a four-day measurement at Mt. Wilson over the Los Angeles basin in January, 1981. We have been unable to reoccupy this site for an actual strain measurement.

The present study, however, has given us some very encouraging results. Even with no atmospheric correction at all, the range ratios could be determined to an accuracy of one part in  $10^7$ . This is sufficient for an order-of-magnitude estimate of incremental shear strain in the southern California region if two measurements separated by one to two years are available. Higher accuracies would be attainable with repeated measurements.

In this report, we first describe the advantages and problems of ground-to-ground ranging by a TLRS, leading to the use of relative lateration, or range-ratio method, and its relationship to the regional strain (Section II). Then, we present the data and analysis of the Mt. Wilson experiment (Section III). This is followed by a short treatment of regional strain determination using hypothetical data (Section IV). Finally, we present the conclusions from this feasibility study and offer some recommendations. Some pertinent data are presented in the Appendix.

## II. TLRS GROUND-TO-GROUND RANGING

### Advantages and Problems

The TLRS is a highly mobile satellite laser ranging system designed to perform ground-to-LAGEOS range measurements. It is also highly sensitive, being capable of determining the range to a LAGEOS satellite with return signals as low as one photoelectron every 20 to 50 laser shots [Silverberg et al. 1982]. Used as a ground-to-ground ranging device, it can measure the distance to any single 1 inch (25 mm) corner reflector within sight at very low laser power level. The measureable range is limited only by the curvature of the earth. The required power level is so low that, unlike some systems used for similar measurements, the laser beam can be maintained many orders of magnitude below the eye-damage threshold.

The practical precision of the TLRS range measurements is limited to about 1.5 cm for a one minute average, which is somewhat worse than those of conventional EDM devices using modulated laser beams. However, the long range capability of the TLRS reduces the relative error to well within the limits of interest in conventional surveys. The TLRS has an automatic pointing system, an automatic calibration system and other features which lend themselves to providing many horizontal (ground-to-ground) line measurements on an operational basis. Thus it will be a good device to use if the data it provides is sufficient to determine the regional deformation at a high-enough accuracy.

The most serious problem in using the TLRS for ground-to-ground ranging is the atmospheric effect. The temperature, pressure, and to a lesser degree water vapor influence the index of refraction of air, and thus the speed of a laser beam through the atmosphere. To obtain the absolute distance between two points from a time of flight measurement through air, one must make corrections for these atmospheric variables.

Estimates of these atmospheric variables along the beam path may be made based on measurements at the two end points. This, however, is unsatisfactory for long lines. A more precise way is to measure directly the atmospheric condition along the beam path by flying an aircraft during the ranging. This, though done in practice, is a costly operation. A third alternative is to use more than one wavelength for ranging. Using the dispersive characteristics of light in air, one can correct for the atmospheric effects [Huggett, et al. 1977].

The present TLRS operates in a single color. Flying an aircraft, we judged, is too costly for repeated measurements in many directions. Thus we had to look for another alternative.



### Relative Lateration

One way to improve the accuracy of range measurements without relying on expensive in-flight measurement of atmospheric conditions is to use a relative lateration technique, or the "ratio method" (Robertson, 1972). Instead of attempting to measure the absolute length of each survey line to high accuracy, this technique determines only the ratios of distances. This method is based on a supposition that the temporal changes of atmospheric conditions along several survey lines within a given region are similar to each other. Therefore, even when the time of flight of a laser beam in each line fluctuates with changing atmospheric conditions, the ratios of the times of flight along different survey lines tend to vary little with time.

Carter and Vincenty [1978] used this method in an experimental EDM survey around the McDonald Observatory in 1977. They obtained sets of measurements, one month apart, consistent to one to two parts in  $10^7$ . They have just repeated this experiment and the data is now being analyzed. Since the results of Carter and Vincenty appear to be quite promising, we have decided to try the same for our TLRS measurements.

### Relationship Between Relative Lateration and Strain

Unlike absolute measurements of distances, the relative distance measurements repeated after a certain time period will not give all of the components of deformation, or incremental strain, for the time period unless at least one survey line is measured absolutely. However, a clear relationship exists between the changes of relative distances and incremental shear strain.

Let us consider  $n$  survey lines radiating from a central station. In the present case, the TLRS is located at the central station and a retroreflector is located at the end of each radiating line. Assume that all lines lie in a horizontal plane, neglecting both the curvature of the earth's surface and topographic height differences. Choosing a coordinate system with the origin at the central station, positive  $x$  towards east and positive  $y$  towards north, the original length of line  $i$  to the reflector at coordinates  $(x_i, y_i)$  at the time of the initial survey is given by

$$s_i = (x_i^2 + y_i^2)^{1/2} \quad (1)$$

Now assume that between the initial survey and a subsequent survey the entire area of the survey undergoes a uniform deformation represented by incremental strain components  $\epsilon_{xx}$ ,  $\epsilon_{xy}$  and  $\epsilon_{yy}$ . Then, the line length becomes

$$\begin{aligned}
 s_i' &= [(x_i + \epsilon_{xx}x_i + \epsilon_{xy}y_i)^2 + (y_i + \epsilon_{xy}x_i + \epsilon_{yy}y_i)^2]^{1/2} \\
 &= [x_i^2 + y_i^2 + 2\epsilon_{xx}x_i^2 + 4\epsilon_{xy}x_iy_i + 2\epsilon_{yy}y_i^2]^{1/2} \\
 &= s_i [1 + 2\epsilon_{xx}\sin^2\alpha_i + 4\epsilon_{xy}\sin\alpha_i \cos\alpha_i + 2\epsilon_{yy}\cos^2\alpha_i]^{1/2} \\
 &= s_i [1 + \epsilon_{xx}\sin^2\alpha_i + \epsilon_{xy}\sin 2\alpha_i + \epsilon_{yy}\cos^2\alpha_i] \quad (2)
 \end{aligned}$$

where  $\alpha_i = \tan^{-1}(x_i/y_i)$  is the azimuth of the line  $i$  measured clockwise from north, and the higher order terms in strain have been neglected. Then, the range increment  $\delta_i$  is given by

$$\delta_i = s_i' - s_i = s_i [\epsilon_{xx}\sin^2\alpha_i + \epsilon_{xy}\sin 2\alpha_i + \epsilon_{yy}\cos^2\alpha_i] \quad (3)$$

Next define original mean range and range ratios to the mean, respectively, as

$$\bar{s} = \frac{1}{n} \sum_{i=1}^n s_i \quad (4)$$

and

$$r_i = s_i / \bar{s} \quad (5)$$

Then, the subsequent mean range and range ratios are

$$\bar{s}' = \frac{1}{n} \sum_{i=1}^n s_i' / n = \bar{s} + \frac{1}{n} \sum_{i=1}^n \delta_i \quad (6)$$

and

$$\begin{aligned}
 r_i' &= s_i' / \bar{s}' = (s_i + \delta_i) / (\bar{s} + \frac{1}{n} \sum_{i=1}^n \delta_i) \\
 &= r_i (1 + \delta_i / s_i - \frac{1}{n} \sum_{i=1}^n \delta_i / \bar{s}) \quad (7)
 \end{aligned}$$

where the higher order terms are again neglected. The increment of the range ratio is, therefore,

$$r_i' - r_i = r_i (\delta_i / s_i - \frac{1}{n} \sum_{i=1}^n \delta_i / \bar{s}) \quad (8)$$

Then, range ratio increment normalized by the original range ratio is given by

$$\gamma_i = (r_i' - r_i)/r_i = \delta_i/s_i - \frac{1}{n} \sum_{i=1}^n \delta_i / \bar{s} \quad (9)$$

Substituting eq. (3) into eq. (9), and using (5), we obtain

$$\begin{aligned} \gamma_i = & [\sin^2 \alpha_i - \frac{1}{n} \sum_{i=1}^n (r_i \sin^2 \alpha_i) / n] \epsilon_{yy} \\ & + [\sin 2\alpha_i - \frac{1}{n} \sum_{i=1}^n (r_i \sin 2\alpha_i) / n] \epsilon_{xy} \\ & + [\cos^2 \alpha_i - \frac{1}{n} \sum_{i=1}^n (r_i \cos^2 \alpha_i) / n] \epsilon_{yy} \end{aligned} \quad (10)$$

Equation (10) may give one an impression that a set of measurements of the normalized range ratio increments  $\gamma_i$  would give the incremental strain components  $\epsilon_{xx}$ ,  $\epsilon_{xy}$  and  $\epsilon_{yy}$ . However, this impression is incorrect because the coefficients of  $\epsilon_{xx}$  and  $\epsilon_{yy}$  are not independent of each other, as their sum vanishes, and therefore  $\epsilon_{xx}$  and  $\epsilon_{yy}$  cannot be determined uniquely.

Now let

$$\Theta = \epsilon_{xx} + \epsilon_{yy} \quad (11)$$

and

$$\Psi = \epsilon_{xx} - \epsilon_{yy} \quad (12)$$

Then,

$$\epsilon_{xx} = \frac{1}{2} (\Theta + \Psi) \quad (13)$$

and

$$\epsilon_{yy} = \frac{1}{2} (\Theta - \Psi) \quad (14)$$

Substituting (13) and (14) into (10), we obtain

$$\begin{aligned} \gamma_i = & [\sin 2\alpha_i - \frac{1}{n} \sum_{i=1}^n (r_i \sin 2\alpha_i) / n] \epsilon_{xy} \\ & - \frac{1}{2} [\cos 2\alpha_i - \frac{1}{n} \sum_{i=1}^n (r_i \cos 2\alpha_i) / n] \Psi \end{aligned} \quad (15)$$

The coefficients of  $\epsilon_{xy}$  and  $\Psi$  are known quantities for the initial setup of the survey lines. Thus, for a set of measurements of the normalized range ratio increments  $\gamma_i$ , the incremental shear strain components  $\epsilon_{xy}$  and  $\Psi$  can be determined by a least-square inversion of eq. (15).

Finally, the maximum incremental shear strain  $S$  and the direction of the principal strain axes  $\beta$  are given by

$$S = [(2\epsilon_{xy})^2 + \Psi^2]^{1/2} \quad (16)$$

and

$$\beta = \frac{1}{2} \tan^{-1} (2\epsilon_{xy} / \Psi) \quad (17)$$

The dilation  $\Theta$  of eq. (11) disappears from eq. (15), and thus cannot be determined. This is expected because any uniform compression or expansion of the entire area causes no change in range ratios.

The treatment above assumes uniform deformation of the entire region. If for some reason, such as the existence of active faults within the area, the regional deformation is not uniform, large residuals will show up in the least-square inversion of eq. (15). Thus, any residuals significantly larger than the measurement errors will indicate heterogeneous strain.

The remaining question is how accurately we can estimate the normalized range ratio increments  $\gamma_i$ . Since the measurements are done in terms of time of flight of light beams, the uncertainty in speed of light is the determining factor. The average speed of light,  $c_i$ , between the central station and a reflector  $i$  may be expressed as the sum of four components:

$$c_i = c_0 + \ell_i + w_c + w_i \quad (18)$$

where  $c_0$  is the speed of light in standard air, which is constant for all survey lines at all times;  $\ell_i$  is the correction attributable to the reflector location, which is time invariant for a given reflector;  $w_c$  is a component of correction attributable to weather common to all reflectors at a given time; and  $w_i$  is the residual weather correction. The line length  $s_i$  is given in terms of round-trip time of flight,  $t_i$ , as

$$s_i = \frac{1}{2}(c_0 + \ell_i + w_c + w_i)t_i \quad (19)$$

the mean range as

$$\bar{s} = \frac{1}{2}(c_0 \bar{t} + \sum_{i=1}^n \ell_i t_i / n + w_c \bar{t} + \sum_{i=1}^n w_i t_i / n) \quad (20)$$

where  $\bar{t} = \sum_{i=1}^n t_i / n$  is the mean time of flight, and the range ratio to the mean as

$$r_i = u_i [1 + (\ell_i - \sum_{i=1}^n \ell_i t_i / n \bar{t} + w_i - \sum_{i=1}^n w_i t_i / n \bar{t}) / c_0] \quad (21)$$

where  $u_i = t_i / \bar{t}$  is the time-of-flight ratio to the mean, and the higher-order terms have been neglected. Finally, the normalized range ratio increment is given as

$$\gamma_i = \eta_i + [(w_i' - w_i) - \sum_{i=1}^n (w_i' - w_i) t_i / n \bar{t}] / c_0 \quad (22)$$

where  $\eta_i = (u_i' - u_i)/u_i$  is the normalized time-of-flight ratio increment and quantities with primes designate those at subsequent measurement as before. The higher-order terms are again neglected. Note that the common weather component,  $w_c$ , is eliminated by taking the range ratio (21), and the location specific components,  $\ell_i$ 's, are eliminated by normalization (22), leaving only the residual weather components  $w_i$ 's.

The normalized time-of-flight ratio increment,  $\eta_i$ , thus approximates the range ratio increments,  $\gamma_i$ , with a small error due to residual weather term. The latter is not location specific, and is not common to all lines at a given time. If this term is sufficiently small, then we can substitute  $\eta_i$  for  $\gamma_i$  in calculating the shear strain increment using (15).

### III. MT. WILSON EXPERIMENT

#### Field Experiment

At the request of the NASA Crustal Dynamics Project, the TLRS team from the McDonald Observatory of the University of Texas, led by Dr. Eric Silverberg, deployed the TLRS at Mt. Wilson, California, in January of 1981. At the same time, two of us (H.J.D. and T.C.) scouted the surrounding area for suitable target sites and selected the reflector locations. Then, a field party from the National Geodetic Survey (NGS), which was dispatched at the request of NASA to help us, deployed retroreflectors at the chosen sites. The survey lines selected for the site are shown in Figure 1. Table 1 lists the nominal coordinates of the base station (TLRS site) at Mt. Wilson and of the end points of the lines, where the retroreflectors were installed. Also listed in Table 1 are the approximate look angles from Mt. Wilson and ranges as computed from the indicated coordinates using the IAG standard ellipsoid Geodetic Reference System 1967.

Each reflector except the one at Cahuenga was a metal box containing an array of three 1½ inch (38mm) corner cubes, supplied by the NGS. The box was mounted on a tripod and placed directly over the station mark using an optical blumb bob. This elaborate configuration made it necessary to guard the reflector continuously for the entire duration of the experiment. The reflector used at Cahuenga was designed by one of us (T.C.) for unmanned operation. It contained a single 1 inch (25mm) corner cube and was fastened to an outcrop with anchor bolts at a site off the station mark, thus concealed from public view.

The reference point of the TLRS, from which the raw time-of-flight measurements were made, was slightly offset from the Mt. Wilson station mark given in Table 1. The measured coordinates of the station mark relative to the TLRS were:

$$x = -1.4873 \text{ m (west)}$$

$$y = 0.5093 \text{ m (north)}$$

$$z = -3.3709 \text{ m (below)}$$

The resulting corrections, to be applied to the observed quantities to reduce them to the reference mark, are listed in Table 2. The corrections can be applied at any stage of data reduction.

After the initial setup, which began on January 9, 1981, the horizontal ranging data were collected over the four-day interval January 23 through 26, 1981, in cooperation with the NOAA National Geodetic Survey. Each of the reflector sites except Cahuenga was manned continuously during the entire experiment to record the temperature, pressure and relative humidity at the site at about 30 minute intervals. The details of the data acquisition are given in Silverberg et al. [1982].

1981 MT. WILSON  
RADIAL LINE SURVEY

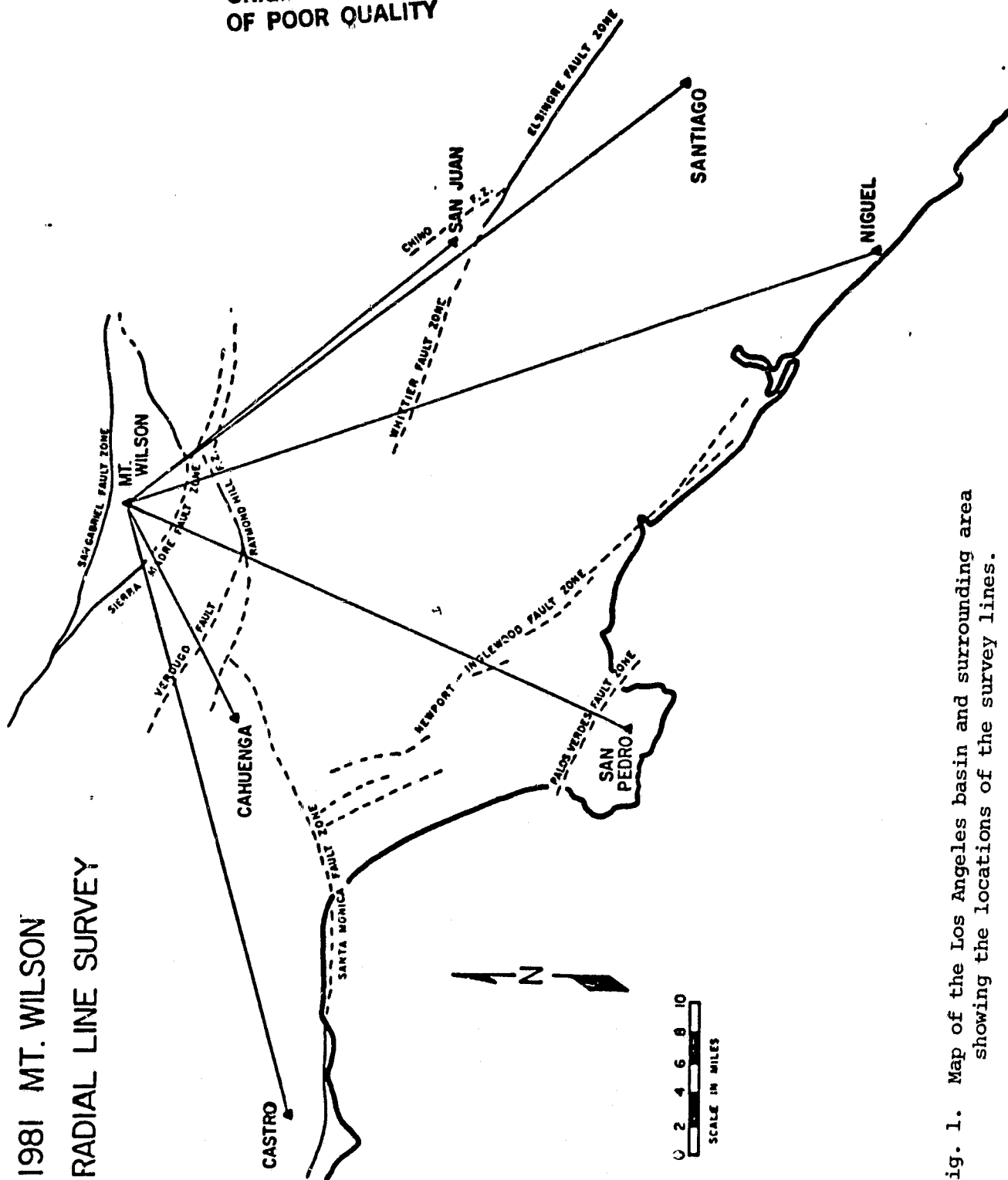


Fig. 1. Map of the Los Angeles basin and surrounding area showing the locations of the survey lines.

ORIGINAL PAGE 12  
OF POOR QUALITY

Table 1. Stations Used in Mt. Wilson Line Survey

Station	Longitude	Latitude	Elevation m	Look Angle		Range m	Note
				Azimuth	Altitude		
Mt. Wilson	241° 56' 17.85"	34° 13' 21.58"	1722.0				1
Castro	241° 12' 55.40"	34° 05' 08.57"	858.9	257.366	-1.030	68393	2
Cahuenga	241° 40' 29.81"	34° 08' 13.00"	554.1	248.687	-2.681	26104	3
San Pedro	241° 39' 55.73"	33° 44' 40.88"	447.1	205.506	-1.508	58729	4
Niguel	242° 16' 00.18"	33° 30' 44.83"	288.0	158.814	-1.353	84459	
Santiago	242° 28' 00.10"	33° 42' 37.89"	1733.0	139.168	-0.329	74933	
San Juan	242° 15' 45.99"	33° 54' 49.47"	543.0	138.751	-1.688	45536	

1 New marker 9.408m from Mt. Wilson E10A @345°17'

2 Solitice Canyon B2 Aux. 1, which is 14.107m FNE of Castro 1898

3 Reference mark #3 of Cahuenga #2, 13.329m @257°47' from Cahuenga #2

4 L7 Ecc. San Pedro Hills, which is 12.576m @318°57' from San Pedro #3

Table 2. Corrections to be Applied to Observations to Reduce to Mt. Wilson Gound Marker

Station	Round-Trip		Range Ratio*	
	Time of Flight ns	Range m	(1) ppm	(2)
Castro	-9.344	-1.4003	-25.08	-26.78
Cahuenga	-9.054	-1.3568	-23.35	-23.51
San Pedro	-1.798	-0.2694	-5.91	-7.93
Niguel	6.222	0.9325	13.61	9.93
Santiago	8.931	1.3385	20.64	17.07
San Juan	8.432	1.2636	20.09	17.64

\* (1) Ratio to mean range

(2) Same but excluding Cahuenga and Niguel



## The Data

The raw field data were initially processed at the University of Texas at Austin by the McDonald Observatory group. As described in detail by Silverberg *et al.* [1982], the processing of the raw data involved accumulation of individual photon returns into 200 psec bins, smoothing of the coadded returns by three-bin (600 psec) running averages, cross-correlation with a reference standard to eliminate long-term drift in the calibration constants, adjustments to account for certain measurement irregularities, and removal of a 86.8 nsec constant calibration correction.

The calibrated round-trip time-of-flight data, shown in Figure 2 and listed in Table A1 in the Appendix, have not been corrected for the offset of the TLRS from the ground marker (Table 2). The data for Cahuenga were not used for the analysis because of certain processing difficulties encountered for the data for this station.

The data gap during the second day of observation was due to an interruption in data acquisition caused by rain which accompanied the passage of a cold front. The meteorological data taken at Mt. Wilson site and other stations are shown in Figures A1 through A3, and are listed in Table A2 in the Appendix.

As expected, the raw time-of-flight data show large fluctuations, which are only partially correlated with the meteorological data. The relative RMS deviations of the time-of-flight data (Table 3, column 3) range from 1.53 ppm for Niguel, which was surveyed only after the passage of the cold front, to 3.78 ppm for San Juan, which was the shortest line. The weighted average for all lines is 2.84 ppm.

## Range Ratios to a Single Reference Line

Silverberg *et al.* [1982] calculated the time-of-flight ratios and atmosphere-corrected range ratios to a reference line following the procedure used by Carter and Vincenty [1978]. The reference line they chose was a smoothed curve (a cubic spline) through the Santiago data. Their results (Table 3, column 5) show relative RMS deviations of time-of-flight ratios ranging from 0.4 ppm for Niguel to 1.6 ppm for San Juan. The weighted average for all lines is 1.0 ppm, which is about a factor of three improvement from the fluctuation of the time-of-flight data.

Their results for the range ratios with atmospheric corrections based on end-point meteorological data did not fare as well. In fact the relative RMS deviations increased typically about 40% from those of uncorrected time-of-flight ratios [Silverberg *et al.* 1982].

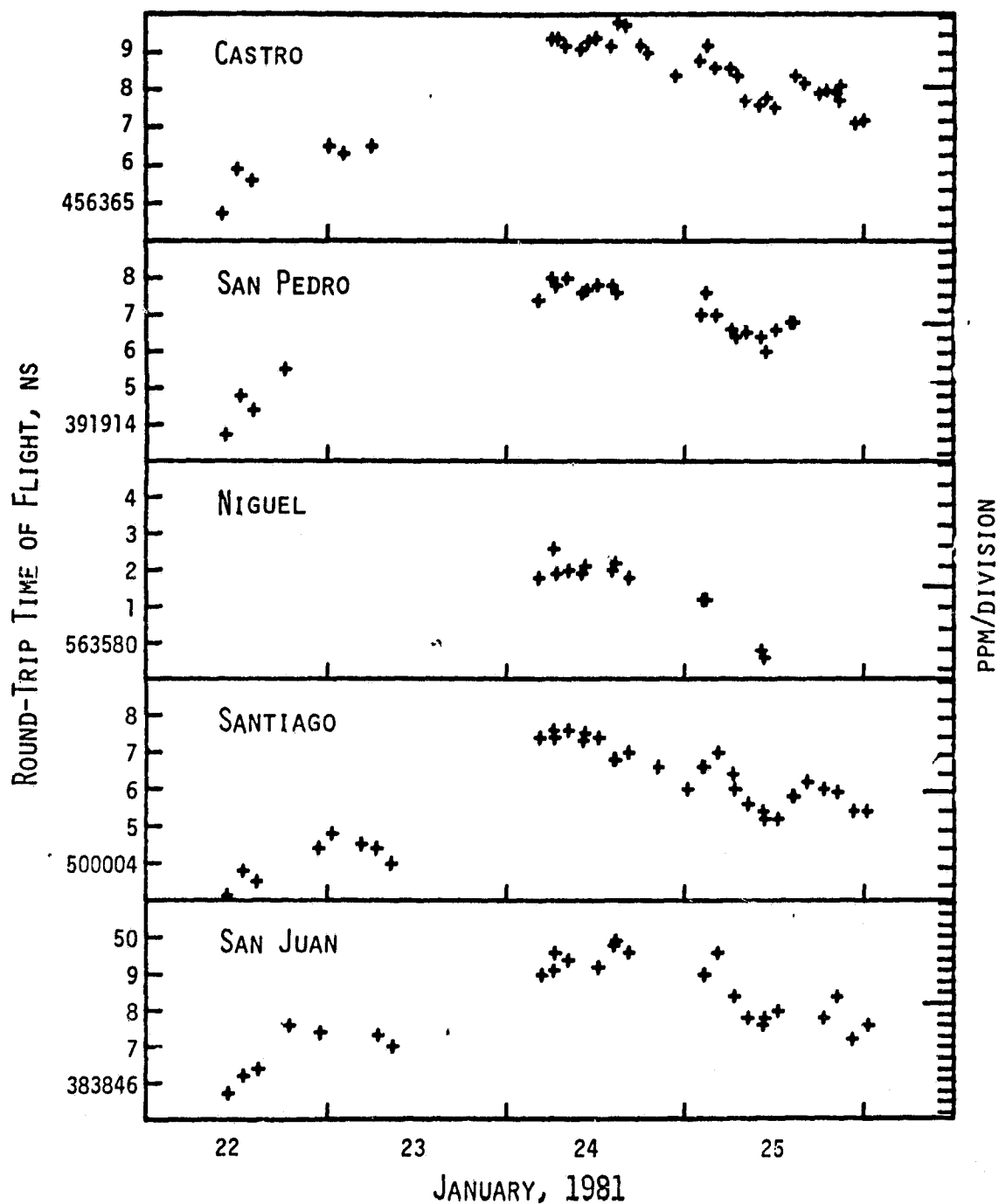


Fig. 2. Round-trip time of flight from Mt. Wilson. The data are not corrected for the marker offset.

The main reason for the poor performance of atmosphere-corrected values is the difficulty of making proper atmospheric corrections. A comparison of the variations of the group index of refraction calculated from the temperature and pressure at end points (Figure 3; also listed in Table A3 in the Appendix) with the time-of-flight variations (Figure 2) clearly shows that long-term variations are fairly well matched but shorter diurnal fluctuations are larger for the index of refractions than for the times-of-flight. Thus the index-of-refraction correction per Carter and Vincenty [1978] over-compensates for diurnal variations.

#### Time-of-Flight Ratios to the Mean

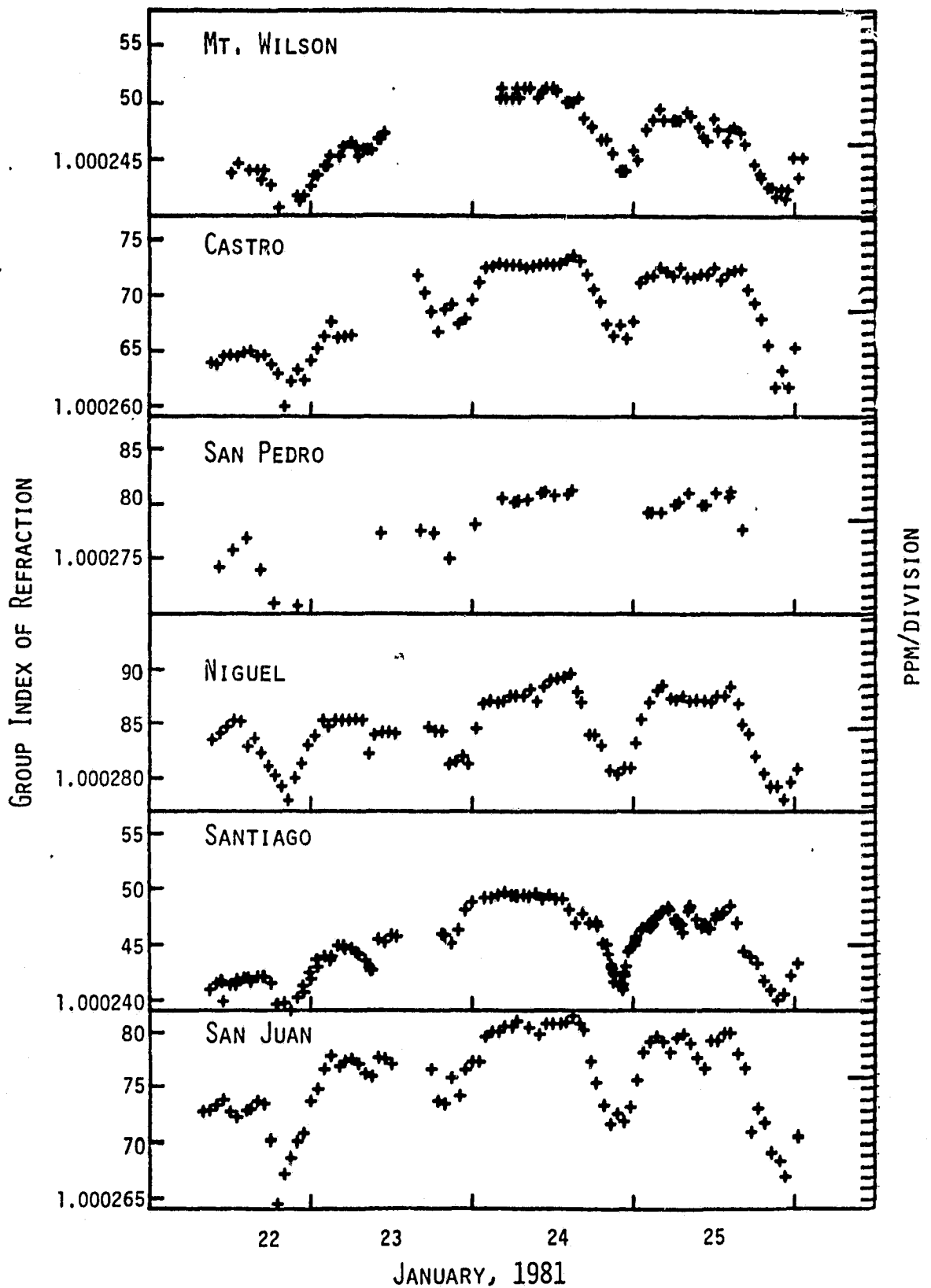
In order to be consistent with the range-ratio/strain relationship of the preceding section, we calculated the time-of-flight ratios to the mean. Since the time-of-flight measurements to all targets were not made exactly simultaneously, the data were linearly interpolated before the mean time-of-flight for a given time was calculated. (Higher order interpolations or a spline approximation might be better, but we judged the difference would be small.) Also, since we had data for Niguel only during the last half of the experiment, this station was excluded from the mean time-of-flight calculation.

The resulting time-of-flight ratios to the mean (Figure 4; also listed in Table A4 in the Appendix) show a further improvement in the fluctuations of the results. The relative RMS deviations (Table 3, column 5) now range from 0.36 ppm for Niguel to 1.24 ppm for San Juan, with the weighted average of 0.71 ppm for all lines, a factor of four improvement from the raw time-of-flight data.

#### An Alternative Atmospheric Correction

As stated earlier, the short-term, diurnal fluctuations in the index of refraction at end points exceed the observed fluctuations in the time-of-flight values. This is probably due to the larger fluctuation of the atmospheric temperature near the ground than those in most of the intervening air mass; a result of the base station and most of the target stations being located well above the intervening terrain. In this situation, a standard correction procedure like that of Carter and Vincenty [1978] is not really applicable, and some alternate procedures are needed.

An experimental procedure we tried was to estimate the average temperature of the air mass by low-pass filtering the mean of the temperatures measured at the end points. The filter we used was a simple one of adding all previous temperature readings each weighted by a factor proportional to a negative exponential of the elapsed time. After



Fgi. 3. Group index of refraction computed from atmospheric data at end points.

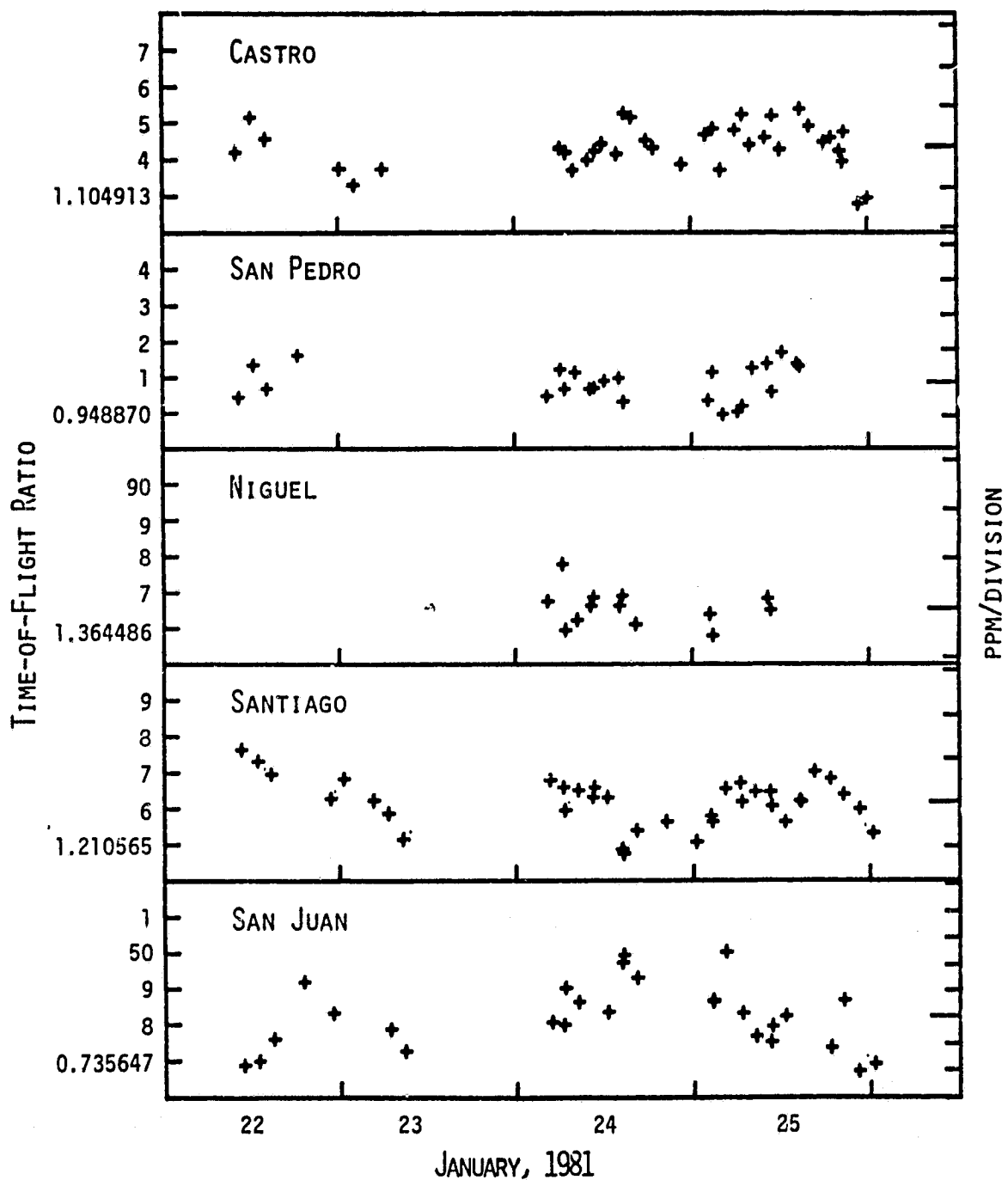


Fig. 4. Time-of-flight ratios to the mean.

trials with such filters of several different time constants, a time constant of 12 hours was found to give the best result. The resulting relative RMS deviations of the ranges thus corrected for atmospheric conditions (Table 3, column 4) range from 0.96 ppm for Castro to 2.05 ppm for San Juan with the weighted average of 1.38 ppm for all lines. This is about a factor of two improvement from the raw time-of-flight data. However, the range ratios calculated from these corrected ranges do not show any significant improvement over those of the uncorrected ratios. The relative RMS deviations of the corrected range ratios (Table 3, last column) range from 0.52 ppm for Niguel and Santiago to 1.21 ppm for San Juan, with the weighted average of 0.72 ppm for all lines.

A comparison of the relative RMS deviations of various quantities in Table 3 reveals that the best result is obtained for the uncorrected time-of-flight ratios to the mean. The atmospheric corrections did not improve the RMS deviations at all when ratios were taken.

#### A Test for Systematic Error Due to Atmospheric Conditions

The reliability of the relative lateration depends on the validity of the assumption that the temporal changes of atmospheric conditions are similar for all survey lines in the area so that their effects cancel out when ratios are taken. If this assumption is incorrect, a systematic error due to varying atmospheric conditions is introduced into the measured time-of-flight ratios. The greatly different atmospheric conditions before and after the passage of a cold front during the experiment gave us an opportunity to test this assumption.

The test we performed is the likelihood ratio test. We divided the time-of-flight ratios of Table A4 for each line into two subsets, the first half and the last half, of equal size (the last half was one greater than the first half if the total number was odd). If the systematic error due to atmospheric conditions is significantly large, the mean ratio,  $\mu_1$ , for the first subset will be significantly different from that,  $\mu_2$ , for the second subset. Setting up a null hypothesis  $H_0: \mu_1 = \mu_2$ , if it is true, then the likelihood ratio statistic

$$t = [n_1 n_2 / (n_1 + n_2)]^{1/2} (\mu_1 - \mu_2) / [(n_1 \sigma_1^2 + n_2 \sigma_2^2) / (n_1 + n_2 - 2)]^{1/2} \quad (23)$$

has a t distribution with  $n_1 + n_2 - 2$  degrees of freedom, where  $n_1$  and  $n_2$  are the sample sizes of the two subsets and  $\mu_1, \mu_2, \sigma_1^2$  and  $\sigma_2^2$  are used to designate the sample means and the sample variances of the first and the second subsets, respectively, for convenience.

At 90% significance level, the t distribution has values of 1.69 for 34 degrees of freedom and 1.80 for 11 degrees of freedom, while the t values computed from the data, Table 4, are much smaller. Therefore, the null hypothesis cannot be rejected at this level of significance.

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Table 3. Comparison of Relative RMS Deviations in ppm

Station	Number of Data Points	Uncorrected Time of Flight	Corrected Range(a)	Uncorrected T-o-F Ratio to Santiago(b)	Uncorrected T-o-F Ratio to the mean	Corrected Range Ratio to the mean
Castro	36	2.68	0.96	0.7	0.56	0.54
San Pedro	24	2.91	1.12	0.8	0.52	0.54
Niguel	13	1.53	1.81	0.4	0.36	0.52
Santiago	36	2.45	1.03	(0.2)	0.54	0.52
San Juan	27	3.78	2.05	1.6	1.24	1.21
All	136	2.84	1.38	1.0	0.71	0.72

(a) Corrected by using 12-hour low-pass filtered temperature

(b) From Silverberg et al. [1982]. The deviation for Santiago is from the smoothed curve, and is not included in calculating the average for all stations.

Table 4. Likelihood Ratio Test for Non-equality of Means

Station	Subset 1			Subset 2			Degree of Freedom	t
	$n_1$	$\mu_1$	$\sigma_1$	$n_2$	$\mu_2$	$\sigma_2$		
Castro	18	1.104914282	0.000000513	18	1.104914438	0.000000694	34	0.745
San Pedro	12	0.948870911	0.000000356	12	0.948870818	0.000000598	22	0.445
Niguel	6	1.364486713	0.000000572	7	1.364486460	0.000000365	11	0.889
Santiago	18	1.210566249	0.000000774	18	1.210566130	0.000000515	34	0.531
San Juan	13	0.735648159	0.000000830	14	0.735648288	0.000000983	25	0.353

In other words, no significant difference is found between the mean time-of-flight ratios in the first and second halves of the experiment for any of the lines surveyed.

## Results

Since there is no evidence for systematic errors caused by atmospheric conditions, the most likely estimates of the mean time-of-flight ratios and their variances (and standard deviations) can be calculated from the entire data set. The results are shown in Table 5. Also listed in this table are the mean time-of-flight and the mean distances. The latter were calculated using atmospheric corrections based on the low-pass filtered temperatures described earlier and pressures interpolated to the average height of the beam from the end-point measurements (extrapolation in case of Santiago because the average height of the beam was lower than either end point). A group index of refraction of  $n = 1.00028975$  at the wavelength of  $0.5320 \mu\text{m}$ , calculated from the formula<sup>9</sup> given in American Institute of Physics Handbook [1972, p. 6-111] for standard dry air with 0.03% carbon dioxide at  $15^\circ\text{C}$  and 760 mm Hg, is used. No other corrections have been applied to the calculated distances; thus they are subject to minor systematic errors.

The estimated relative standard deviations of the mean time-of-flight ratios are approximately  $1 \times 10^{-7}$  except for San Juan, which is the shortest line. In comparison, Savage and Prescott [1973] estimate that the standard deviation of their Geodolite measurements of distances are 3 and 8 mm for lengths of 1 and 37 km, respectively. Thus, the precision of the present time-of-flight ratios is at least a factor of two better than that of their distance measurements. Furthermore, their distances had to be corrected for temperature and humidity readings made with an aircraft flying along the line of sight, while the present time-of-flight ratios required no atmospheric correction at all.

Multiwavelength measurements of distances are definitely better than the above two in terms of relative accuracy. Huggett and Slater [1975] and Slater and Huggett [1976] show the standard deviation of individual distance measurements to be less than  $1 \times 10^{-7}$  on a 10.1 km line. By taking the mean of many measurements, which is practical in this case, the accuracy can be improved further. The ranges attainable with the multiwavelength system, however, are quite limited compared with the TLRS measurements.



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Table 5. Mean Time of Flight, Distance and Ratios\*

Station	Mean Time of Flight ns	Mean Distance m	Mean T-o-F Ratio	S.D. of Mean T-o-F Ratio	Relative S.D. ppm
Castro	456358.75	68388.52	1.10488758	0.00000010	0.09
San Pedro	391914.94	58730.88	0.94886293	0.00000010	0.11
Niguel	563587.77	84456.72	1.36449651	0.00000014	0.10
Santiago	500014.83	74931.62	1.21058326	0.00000011	0.09
San Juan	303856.63	45535.87	0.77566587	0.00000018	0.24

\* These results have been corrected for the TLRS/ground-marker offset.

Table 6. Increment in Normalized Ratio  
Due to Hypothetical Strain Increment  
and Rounded Values for Testing

Station	Normalized Ratio Increment ppm	Rounded to 0.1 ppm
Castro	0.052	0.1
San Pedro	-0.158	-0.2
Niguel	-0.030	0.0
Santiago	0.068	0.1
San Juan	0.070	0.1

#### IV. SHEAR STRAIN DETERMINATION USING HYPOTHETICAL DATA

We have been unable to reoccupy the Mt. Wilson site for a repeat measurement, which would allow a testing of the ratio method for a shear strain determination in the region. This section, therefore, describes an exercise we have conducted to see how well we can determine the regional shear-strain increment using a set of hypothetical data.

We assume a hypothetical strain increment described by

$$\begin{aligned} \epsilon_1 &= 0.1 \times 10^{-6} & : & \text{maximum extension} \\ \epsilon_2 &= -0.2 \times 10^{-6} & : & \text{maximum compression} \\ \epsilon_1 - \epsilon_2 &= 0.3 \times 10^{-6} & : & \text{maximum shear} \\ \beta &= 110^\circ & : & \text{azimuth of maximum positive principal} \\ & & & \text{axis (extension) measured clockwise} \\ & & & \text{from north} \end{aligned}$$

This strain increment is approximately the annual strain increment in southern California observed by Savage *et al.* [1981]. The resulting increments in normalized range ratios for the five survey lines used in the Mt. Wilson experiment are listed in the center column of Table 6. Since we will not be able to measure these ratio increments at this accuracy, we use the values rounded to  $1 \times 10^{-7}$ , as given in the rightmost column of Table 6.

Substituting these rounded ratio increments into eq. (15), and inverting it in a least-squares sense, we obtain the following results:

$$\begin{aligned} \epsilon_1 - \epsilon_2 &= 0.40 \times 10^{-6} \\ \beta &= 109.5^\circ \end{aligned}$$

The result describes the original hypothetical shear-strain increment reasonably well. A trial with a rounding to  $1 \times 10^{-8}$  results in almost complete duplication of the hypothetical strain increment.

The likelihood ratio test of the preceding section can be used to estimate the required number of measurements to achieve a given level of accuracy at a given confidence level. We use the standard deviation of individual range ratio measurements of  $5 \times 10^{-7}$  as estimated from the present data (Table 3, excluding San Juan). Thus, substituting  $\sigma_1 = \sigma_2 = 10^{-6}$  and  $|\mu_1 - \mu_2| = 10^{-7}$  into eq. (18), we find that  $n_1 = n_2 = 200$  will give  $t = 1.99$ , which exceeds the value of  $t$  distribution, 1.97, for 198 degrees of freedom at 95% confidence level. Thus a variation in the range ratio of  $10^{-7}$  found by averaging 200 ratio measurements is significant at 95% level of confidence.

At a rate of one measurement every hour, it will take slightly more than a week to complete this many measurements. Two such series of

measurements one year apart is sufficient to determine the shear strain increment in southern California.

For a given set of  $t$  and  $\sigma$ 's,  $n$ 's are approximately inversely proportional to the square of the difference in  $\mu$ 's in equation (18). Thus doubling the measurement interval, thereby doubling the expected ratio variations, approximately quarters the required number of measurements. For example, a pair of 50-measurement sets two years apart will give the shear strain rate in southern California.

## V. CONCLUSIONS AND RECOMMENDATIONS

### Conclusions

Even though the field experiment we performed during this contract was quite limited compared with our original plan, we obtained several interesting and important results. The following is a list of conclusions drawn from these results:

1. The increment of the ratio of the length of a survey line to the average of several survey lines in a region is directly related to the incremental shear strain in the region. Thus, the shear strain rate can be calculated from observations of temporal variations in such ratios.
2. Using the TLRS, the time-of-flight ratios could be determined to an accuracy (one standard deviation) of  $1 \times 10^{-7}$  by averaging measurements over a four day period. This accuracy was obtained without using any atmospheric corrections at all. No improvement was obtained when atmospheric corrections based on end-point measurements were applied.
3. A calculation using a hypothetical data simulating the observed strain field in southern California indicates that two sets of TLRS ratio measurements separated by one to two years will be sufficient to determine the direction and rate of shear strain in the region.
4. Thus relative lateration using the TLRS has been demonstrated to be a good method for monitoring the regional shear strain field around satellite ranging stations. The TLRS operates successfully over long distances. The ratio method is extremely economical. It requires no environmental measurements and can be performed with small unattended retroreflectors distributed over a wide area. Thus these techniques greatly surpass the capability of conventional EDM techniques.

### Recommendations

1. The results of the present experiment are thus very encouraging. However, they are based on only one experiment. Before this technique is put to a practical use, further demonstration is needed to confirm the above results. Therefore, it is recommended that this feasibility study be continued at least to include reoccupation of the Mt. Wilson site and two measurements at another properly selected site, preferably with a different meteorological environment.

2. Relative lateration is not limited to the data taken by the TLRS. The data reduction procedure used in the present study can be applied to other data from distance measurements. Therefore, it is recommended that we reanalyze some of existing ranging data to see if improvements in determination of shear strain rate can be achieved. This can be done without further field measurements.
3. Additional feasibility test measurements similar to the Mt. Wilson experiment may be obtained from fixed satellite ranging stations. It is therefore, recommended that this possibility be examined.
4. Horizontal ranging to distant targets on the ground does not require all the sophistication of the TLRS system. Therefore, when the capability of the present technique is fully demonstrated, a smaller, more portable single-photon ranging unit should be developed for this purpose.
5. Finally, the technology is advancing in other fields also. Such techniques as miniature interferometer terminals [Counselman and Shapiro, 1979] may someday be more useful in surveys of regional extent. Therefore, development in these other techniques should be reviewed while developing the present technique.

#### Acknowledgements

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## References

- American Institute of Physics Handbook, 3rd ed., McGraw-Hill, New York, 1972.
- Carter, W. E. and T. Vincenty, Survey of the McDonald Observatory radial line scheme by relative lateration technique, NOAA Technical Report NOS 74 NGS 9, 1978.
- Counselman III, C.C. and I.I. Shapiro, Miniature interferometer terminals for earth surveying, Bull. Geod. 53, 139-163, 1973.
- Huggett, G. R. and L. E. Slater, Precision electromagnetic distance-measuring instrument for determining secular strain and fault movement, Tectonophysics, 29, 19-27, 1975.
- Huggett, G. R., L. E. Slater and J. Langbein, Fault slip episodes near Hollister, California: Initial results using a multiwavelength distance-measuring instrument, J. Geophys. Res., 82, 3361-3368, 1977.
- Robertson, K. D., The use of line pairs in trilateration and traverse, Survey Rev., 21, 290-306, 1972.
- Savage, J. C. and W. H. Prescott, Precision of Geodolite distance measurements for determining fault movements, J. Geophys. Res., 78, 6001-6008, 1973.
- Savage, J. C., W. H. Prescott, M. Lisowski and N. E. King, Strain accumulation in southern California, 1973-1980, J. Geophys. Res., 86, 6991-7001, 1981.
- Silverberg, E. C., and D. L. Byrd, A mobile telescope for measuring continental drift, Sky and Telescope, 61, 405-408, 1981.
- Silverberg, E. C., T. Cahill and J. Dorman, Relative lateration across the Los Angeles basin using a satellite laser ranging system, Bull. Geod., (in press), 1982.
- Slater, L. E. and G. R. Huggett, A multiwavelength distance-measuring instrument for geophysical experiments, J. Geophys. Res., 81, 6299-6306, 1976.

**APPENDIX**

Table A1. Calibrated Round-Trip Time of Flight

Castro				San Pedro				Niguel				Santiago				San Juan			
dy	hr	mn	sc	ns	dy	hr	mn	sc	ns	dy	hr	mn	sc	ns	dy	hr	mn	sc	ns
22	10	04	29	456364.7	22	10	18	55	391913.7	24	04	20	32	563581.8	22	10	41	41	500003.1
22	12	02	41	456365.9	22	12	21	52	391914.8	24	06	19	58	563582.6	22	12	43	11	500003.8
22	14	05	43	456365.6	22	14	15	03	391914.4	24	06	46	28	563581.9	22	14	34	32	500003.5
23	00	09	40	456366.5	22	18	23	39	391915.5	24	08	18	26	563582.0	22	18	55	03	303847.6
23	02	09	25	456366.3	24	04	20	40	391917.4	24	10	15	13	563581.9	22	22	54	50	303847.4
23	06	05	03	456366.5	24	06	16	19	391918.0	24	10	40	16	563582.1	23	06	47	01	303847.3
24	06	08	46	456369.4	24	06	51	04	391917.8	24	14	15	10	563582.0	23	08	47	14	303847.0
24	06	58	17	456369.4	24	08	09	40	391918.0	24	14	40	18	563582.2	24	04	40	14	303849.0
24	08	00	19	456369.2	24	10	10	03	391917.6	24	16	18	20	563581.8	24	06	29	45	303849.1
24	09	59	54	456369.1	24	10	45	34	391917.7	25	02	21	58	563581.2	24	06	35	37	303849.6
24	10	55	03	456369.3	24	12	10	04	391917.8	25	02	48	49	563581.2	24	08	25	49	303849.4
24	12	00	58	456369.4	24	14	11	30	391917.8	25	10	16	05	563579.8	24	12	25	08	303849.2
24	14	01	34	456369.2	24	14	45	60	391917.6	25	10	41	13	563579.6	24	14	25	20	303849.8
24	14	55	05	456369.8	25	02	13	18	391917.0	24	10	35	15	500007.5	24	14	35	04	303849.9
24	16	01	03	456369.7	25	02	53	27	391917.6	24	12	20	03	500007.4	24	16	28	54	303849.6
24	18	00	06	456369.2	25	04	09	53	391917.0	24	14	20	11	500006.8	25	02	33	03	303849.0
24	18	57	39	456369.0	25	06	10	05	391916.6	24	14	35	11	563806.8	25	02	36	56	303849.0
24	22	44	42	456368.4	25	06	45	43	391916.4	24	16	29	53	500007.0	25	04	25	60	303849.6
25	02	05	25	456368.8	25	08	11	36	391916.5	24	18	04	04	500006.6	25	06	32	42	303849.4
25	02	58	04	456369.2	25	10	10	16	391916.4	24	20	18	04	500006.6	25	08	25	56	303847.8
25	04	00	52	456368.6	25	10	45	23	391916.0	25	00	25	13	500006.0	25	10	25	11	303847.6
25	06	01	17	456368.6	25	12	10	30	391916.6	25	02	29	30	500006.6	25	10	36	03	303847.0
25	06	55	21	456368.4	25	14	15	11	391916.8	25	04	22	37	500007.0	25	12	25	03	303849.0
25	08	04	49	456367.7	25	14	40	42	391916.8	25	06	22	51	500006.4	25	18	51	55	303847.0
25	10	01	53	456367.6						25	08	35	48	500006.0	25	20	25	59	303847.2
25	10	55	08	456367.8						25	08	20	32	500005.6	25	22	28	50	303847.2
25	12	04	05	456367.5						25	10	20	08	500005.4	26	00	32	50	303847.6
25	14	51	54	456368.4						25	10	35	11	500005.2					
25	16	00	28	456368.2						25	12	21	02	500005.2					
25	18	00	19	456367.9						25	14	21	41	500005.8					
25	18	59	43	456368.0						25	14	35	02	500005.8					
25	20	07	42	456367.9						25	16	20	06	500006.2					
25	20	37	24	456367.7						25	18	39	42	500006.0					
25	20	49	58	456368.1						25	20	20	59	500005.9					
25	22	51	57	456367.1						25	22	37	54	500005.4					
26	00	05	27	456367.2						26	00	26	53	500005.4					



Table A2. Temperature, T, Pressure, P, and Saturated Vapor Pressure, V

Time		T		P		V		Time		T		P		V	
dy	hr	mm	sc	deg	C	mmHg	mmHg	dy	hr	mm	sc	deg	C	mmHg	mmHg
Mt. Wilson															
22	12	08	01	8.0	624.1	8.0422		24	13	54	08	0.0	621.8	4.5814	
22	13	08	54	7.0	624.1	7.5104		24	14	28	00	0.0	621.8	4.5814	
22	14	53	40	8.0	624.8	8.0422		24	14	54	16	0.0	621.8	4.5814	
22	15	59	57	8.0	624.8	8.0422		24	15	53	54	0.0	622.6	4.5814	
22	16	31	16	9.0	624.8	8.0422		24	16	41	49	2.0	622.6	5.2917	
22	16	57	27	8.0	624.8	8.0422		24	17	53	26	3.0	623.3	5.6817	
22	17	59	44	10.0	625.6	9.2048		24	19	06	29	4.0	622.6	6.0973	
22	19	14	19	12.0	624.8	10.5136		24	19	55	54	4.0	622.6	6.0973	
22	21	57	23	10.0	623.3	9.2048		24	20	49	38	5.0	621.8	6.5398	
22	22	25	11	11.0	624.1	9.8401		24	21	54	25	6.0	620.3	7.0108	
22	23	02	15	10.0	623.3	9.2048		24	22	27	53	6.0	620.3	7.0108	
23	00	00	56	9.0	623.3	8.0422		24	22	52	51	6.0	620.3	7.0108	
23	00	28	23	8.0	623.3	8.0422		24	23	55	10	4.0	620.3	6.0973	
23	01	03	14	8.0	623.3	8.0422		25	00	34	16	5.0	620.3	6.5398	
23	02	01	37	7.0	623.3	7.5104		25	02	00	11	2.0	620.3	5.2917	
23	02	24	05	7.0	623.3	7.5104		25	03	03	19	1.0	620.3	4.9249	
23	02	48	26	6.0	623.3	7.0108		25	03	55	51	0.0	620.3	4.5814	
23	04	13	00	6.0	623.3	7.0108		25	04	32	31	1.0	620.3	4.9249	
23	04	53	34	5.0	623.3	6.5398		25	05	54	44	1.0	620.3	4.9249	
23	05	57	22	5.0	624.1	6.5398		25	06	28	44	1.0	620.3	4.9249	
23	06	37	07	5.0	623.3	6.5398		25	07	01	06	1.0	620.3	4.9249	
23	06	55	29	6.0	623.3	7.0108		25	07	56	42	0.0	619.6	4.5814	
23	07	52	58	5.0	622.6	6.5398		25	08	36	19	0.0	618.8	4.5814	
23	08	20	32	5.0	622.6	6.5398		25	09	48	16	1.0	618.8	4.9249	
23	08	54	34	5.0	622.6	6.5398		25	10	28	22	2.0	618.8	5.2917	
23	09	56	13	4.0	622.6	6.0973		25	11	00	51	2.0	618.1	5.2917	
23	10	29	17	4.0	622.6	6.0973		25	11	59	30	0.0	618.1	4.5814	
23	10	55	37	3.0	621.8	5.6817		25	12	34	14	1.0	618.1	4.9249	
24	04	06	03	0.0	622.6	4.5814		25	13	57	00	2.0	618.1	5.2917	
24	04	22	39	-1.0	622.6	4.2589		25	14	26	09	1.0	618.1	4.9249	
24	04	55	29	0.0	622.6	4.5814		25	14	56	49	1.0	618.8	4.9249	
24	05	56	14	0.0	622.6	4.5814		25	15	55	10	2.0	619.6	5.2917	
24	06	34	27	-1.0	622.6	4.2589		25	16	36	22	3.0	619.6	5.6817	
24	07	04	10	0.0	622.6	4.5814		25	17	55	39	5.0	619.6	6.5398	
24	07	51	41	-1.0	622.6	4.2589		25	18	45	58	6.0	619.6	7.0108	
24	08	31	26	-1.0	622.6	4.2589		25	19	05	52	6.0	618.8	7.0108	
24	09	46	20	0.0	622.6	4.5814		25	19	55	15	7.0	618.8	7.5104	
24	10	28	48	-1.0	621.8	4.2589		25	20	31	54	7.0	618.8	7.5104	
24	11	00	43	-1.0	622.6	4.2589		25	21	06	50	8.0	618.8	8.0422	
24	11	55	50	-1.0	622.6	4.2589		25	21	54	44	7.0	618.1	7.5104	
24	12	34	12	-1.0	621.8	4.2589		25	22	32	53	8.0	618.1	8.0422	
								25	23	02	07	7.0	618.1	7.5104	
								25	23	48	19	4.0	618.8	6.0973	
								26	00	36	21	6.0	618.8	7.0108	

Castro

22	09	10	00	14.4	690.8	8.0724		22	09	10	00	14.4	690.8	8.0724	
22	10	00	00	14.2	690.1	9.2050		22	10	00	00	14.2	690.1	9.2050	
22	11	00	00	13.5	690.1	8.3828		22	11	00	00	13.5	690.1	8.3828	
22	12	00	00	13.6	690.8	9.2050		22	12	00	00	13.6	690.8	9.2050	
22	13	00	00	13.4	690.1	7.9210		22	13	00	00	13.4	690.1	7.9210	
22	14	00	00	13.3	690.8	7.6258		22	14	00	00	13.3	690.8	7.6258	
22	15	00	00	13.2	693.8	7.7721		22	15	00	00	13.2	693.8	7.7721	
22	16	00	00	13.8	693.8	7.9210		22	16	00	00	13.8	693.8	7.9210	
22	17	00	00	14.0	691.6	7.9210		22	17	00	00	14.0	691.6	7.9210	
22	18	00	00	14.9	691.6	8.2263		22	18	00	00	14.9	691.6	8.2263	
22	19	00	00	15.8	691.6	7.9210		22	19	00	00	15.8	691.6	7.9210	
22	20	00	00	18.8	693.8	9.5532		22	20	00	00	18.8	693.8	9.5532	
22	21	00	00	15.9	690.1	9.5532		22	21	00	00	15.9	690.1	9.5532	
22	22	02	00	14.8	690.1	10.2850		22	22	02	00	14.8	690.1	10.2850	
22	23	02	00	15.8	690.1	11.4747		22	23	02	00	15.8	690.1	11.4747	
23	00	00	00	13.9	690.1	8.5418		23	00	00	00	13.9	690.1	8.5418	
23	01	00	00	12.7	690.1	7.9210		23	01	00	00	12.7	690.1	7.9210	
23	02	00	00	11.8	690.8	7.6258		23	02	00	00	11.8	690.8	7.6258	
23	03	00	00	10.4	690.8	7.9210		23	03	00	00	10.4	690.8	7.9210	
23	04	00	00	11.7	690.1	7.3403		23	04	00	00	11.7	690.1	7.3403	
23	05	00	00	11.5	690.1	7.6258		23	05	00	00	11.5	690.1	7.6258	
23	06	00	00	11.4	690.1	7.6258		23	06	00	00	11.4	690.1	7.6258	
23	16	00	00	6.6	692.3	7.6258		23	16	00	00	6.6	692.3	7.6258	
23	17	00	00	8.5	693.1	8.0724		23	17	00	00	8.5	693.1	8.0724	
23	18	00	00	10.3	693.1	8.8679		23	18	00	00	10.3	693.1	8.8679	
23	19	00	00	12.2	693.1	9.9131		23	19	00	00	12.2	693.1	9.9131	
23	20	00	00	10.1	693.1	8.8679		23	20	00	00	10.1	693.1	8.8679	
23	21	00	00	9.3	692.3	8.5418		23	21	00	00	9.3	692.3	8.5418	
23	22	00	00	11.1	692.3	8.8679		23	22	00	00	11.1	692.3	8.8679	
23	23	00	00	10.6	692.3	8.8679		23	23	00	00	10.6	692.3	8.8679	
24	00	00	00	8.9	692.3	8.2263		24	00	00	00	8.9	692.3	8.2263	
24	01	00	00	7.3	692.3	7.3403		24	01	00	00	7.3	692.3	7.3403	
24	02	00	00	5.9	692.3	6.7978		24	02	00	00	5.9	692.3	6.7978	
24	03	00	00	5.8	692.3	7.0645		24	03	00	00	5.8	692.3	7.0645	
24	04	00	00	5.8	693.1	6.7978		24	04	00	00	5.8	693.1	6.7978	
24	05	00	00	5.9	693.1	6.7978		24	05	00	00	5.9	693.1	6.7978	
24	06	00	00	6.0	693.1	6.5400		24	06	00	00	6.0	693.1	6.5400	
24	07	00	00	6.3	693.8	6.5400		24	07	00	00	6.3	693.8	6.5400	
24	08	00	00	6.5	693.8	6.5400		24	08	00	00	6.5	693.8	6.5400	

Table A2. (continued)

Time		T	P	V	Time		T	P	V	Time		T	P	V
dy	hr mn sc	deg C	mmHg	mmHg	dy	hr mn sc	deg C	mmHg	mmHg	dy	hr mn sc	deg C	mmHg	mmHg
24	09 00 00	6.1	693.1	6.2908	22	12 20 00	13.4	719.3	9.2050	22	19 33 00	18.1	740.3	11.8974
24	10 00 00	5.9	693.1	6.0500	22	14 20 00	12.6	720.1	9.2050	22	20 32 00	19.2	739.6	12.3337
24	11 00 00	5.8	693.1	5.8176	22	16 20 00	15.6	720.1	9.9131	22	21 40 00	16.7	736.8	13.2406
24	12 00 00	6.0	693.1	5.8176	22	18 20 00	19.1	720.8	10.6688	22	22 30 00	15.3	738.8	11.0974
24	13 00 00	5.8	693.1	5.8176	22	22 00 00	18.7	719.3	11.8118	22	23 37 00	13.9	739.6	11.4747
24	14 00 00	5.5	693.1	5.5931	23	10 30 00	11.5	718.6		23	00 38 00	13.1	739.6	11.0653
24	15 00 00	5.1	693.1	5.5931	23	16 30 00	12.1	720.8	9.5532	23	01 45 00	11.6	739.6	11.0653
24	16 00 00	5.6	693.1	5.8176	23	18 20 00	12.4	720.8	10.2850	23	02 32 00	12.3	739.6	10.6688
24	17 00 00	6.8	693.1	6.5400	23	20 30 00	14.7	720.8	11.8974	23	03 35 00	11.6	739.6	10.2850
24	18 00 00	8.2	693.1	6.7978	24	00 30 00	11.3	720.1	9.3777	23	04 40 00	11.7	739.6	10.2850
24	19 00 00	9.3	693.1	7.2012	24	04 20 00	9.1	720.8	8.6708	23	05 30 00	11.6	739.6	10.2850
24	20 00 00	11.2	692.3	7.6258	24	06 10 00	9.4	720.8	8.8679	23	06 32 00	11.2	738.8	9.9131
24	21 00 00	12.0	691.6	7.6258	24	06 45 00	9.3	720.8	8.7361	23	07 30 00	11.3	738.8	9.9131
24	22 00 00	11.0	691.6	8.2263	24	08 10 00	9.0	720.1	7.7721	23	08 30 00	14.1	738.1	9.5532
24	23 00 00	12.0	690.8	8.2263	24	10 10 00	8.3	720.1	8.2263	23	09 29 00	12.7	738.8	10.6688
25	00 00 00	10.3	690.8	7.3403	24	10 45 00	8.2	720.1	7.8610	23	10 35 00	12.4	738.8	10.2850
25	01 00 00	6.6	690.8	6.0500	24	12 10 00	8.6	720.1	7.7721	23	11 30 00	12.4	738.8	10.2850
25	02 00 00	5.8	690.1	6.0500	24	14 10 00	8.4	720.1	7.6258	23	12 40 00	12.5	738.8	10.2850
25	03 00 00	5.7	690.1	6.5400	24	14 45 00	8.1	720.1	7.6258	23	17 40 00	13.1	741.8	11.4747
25	04 00 00	5.3	690.8	5.8176	25	02 10 00	9.8	719.3	8.5418	23	18 34 00	13.5	741.8	13.2486
25	05 00 00	5.7	690.8	5.3762	25	02 52 00	9.8	719.3	8.7036	23	19 30 00	13.5	741.8	11.0653
25	06 00 00	5.7	690.1	5.1669	25	04 10 00	9.8	719.3	8.2263	23	20 40 00	16.3	741.1	12.3337
25	07 00 00	5.0	690.1	5.3762	25	06 10 00	9.4	720.1	6.7978	23	21 33 00	15.0	741.1	11.4747
25	08 00 00	5.6	689.3	5.3762	25	06 45 00	8.6	718.6	6.4143	23	22 30 00	15.5	741.1	11.4747
25	09 00 00	5.5	689.3	5.3762	25	08 10 00	7.8	718.6	6.0500	23	23 26 00	16.2	741.1	11.0653
25	10 00 00	5.3	689.3	5.3762	25	10 10 00	8.8	718.6	6.7978	24	00 30 00	13.0	741.1	9.9131
25	11 00 00	5.3	689.3	5.5931	25	10 45 00	8.8	718.6	6.7978	24	01 40 00	11.0	741.8	9.2350
25	12 00 00	4.7	689.3	5.3762	25	12 10 00	7.8	718.6	6.2908	24	02 40 00	10.7	741.8	9.2050
25	13 00 00	5.5	688.6	5.8176	25	14 10 00	7.8	717.8	7.0645	24	03 50 00	10.8	741.8	9.2050
25	14 00 00	5.2	689.3	5.3762	25	14 22 00	7.6	718.6	6.7978	24	04 35 00	10.7	741.8	9.2050
25	15 00 00	5.3	690.1	5.5931	25	16 10 00	11.4	719.3		24	05 40 00	10.3	741.8	9.2050
25	16 00 00	5.4	690.8	5.3762						24	06 20 00	10.3	741.8	9.8679
25	17 00 00	7.3	690.8	5.9327						24	07 35 00	10.2	741.8	8.8679
25	18 00 00	8.5	690.8	6.2908						24	08 35 00	9.7	741.8	8.8679
25	19 00 00	10.0	690.8	6.6678	22	09 28 00	13.2	738.8	8.8679	24	09 30 00	10.4	741.1	9.2650
25	20 00 00	12.3	690.1	6.7978	22	10 32 00	12.5	738.8	9.2050	24	10 30 00	9.2	741.1	8.5418
25	21 00 00	16.2	689.3	8.2263	22	11 30 00	12.0	738.8	9.2050	24	11 39 00	8.5	741.1	8.2263
25	22 00 00	14.1	688.6	6.5400	22	12 30 00	11.4	738.8	9.2050	24	12 31 00	8.4	741.1	8.2263
25	23 00 00	16.1	689.3	7.3403	22	13 31 00	11.5	738.8	9.7318	24	13 33 00	8.0	740.3	7.9210
26	00 00 00	12.5	690.1	6.7978	22	14 30 00	14.1	739.6	9.9131	24	14 31 00	7.7	740.3	7.6258
					22	15 41 00	13.6	740.3	9.5532	24	15 35 00	9.3	740.3	8.5418
					22	16 36 00	15.0	740.3	10.6688	24	16 17 00	10.6	741.1	9.2050
					22	17 30 00	16.2	740.3	11.0653	24	17 20 00	13.6	741.1	9.9131
					22	18 31 00	17.1	740.3	12.3337	24	18 17 00	13.5	741.1	9.9131

Table A2. (continued)

Time				T deg C	P mmHg	V mmHg	Time				T deg C	P mmHg	V mmHg	Time				T deg C	P mmHg	V mmHg		
dy	hr	mn	sc				dy	hr	mn	sc				dy	hr	mn	sc				dy	hr
24	19	15	00	14.2	740.3	10.2650	22	14	40	00	9.8	623.3	5.3762	24	05	45	00	1.0	622.6	4.3764		
24	20	23	00	16.0	738.8	11.0653	22	15	00	00	10.2	623.3		24	06	25	00	1.1	622.6	4.2774		
24	21	40	00	16.1	738.1	10.6688	22	16	00	00	10.0	624.1	5.7043	24	06	40	00	1.2	622.6			
24	22	30	00	15.5	738.1	10.2850	22	16	45	00	10.0	624.1	5.5931	24	07	40	00	.7	621.8			
24	23	30	00	15.4	738.1	10.2850	22	17	00	00	10.0	624.1		24	08	20	00	.8	621.8	4.5311		
25	00	27	00	13.1	738.1	9.0350	22	18	00	00	10.8	624.1	6.5400	24	09	20	00	.5	621.8	4.4780		
25	01	15	00	11.0	738.1	8.8679	22	18	45	00	13.0	624.1	7.3403	24	10	20	00	.8	621.1	4.1783		
25	02	27	00	9.4	738.1	8.2263	22	19	00	00	13.0	624.1		24	11	20	00	.4	621.1			
25	03	30	00	8.4	738.1	7.6258	22	20	00	00	12.5	623.3		24	12	20	00	.8	621.1	3.8100		
25	04	23	00	7.9	738.1	7.3403	22	21	00	00	13.0	622.6		24	13	20	00	.8	621.1			
25	05	40	00	8.8	737.3	6.7978	22	22	00	00	11.6	622.6	6.2908	24	14	20	00	1.8	621.1	4.5811		
25	06	25	00	8.9	737.3	6.7978	22	22	45	00	10.3	622.6		24	15	20	00	3.2	621.1			
25	07	28	00	8.5	737.3	6.5400	22	23	00	00	11.0	622.6		24	16	20	00	2.2	621.1	4.0943		
25	08	23	00	9.7	736.6	6.7978	22	23	45	00	9.2	623.3	7.0645	24	17	20	00	3.5	621.8			
25	09	29	00	8.3	735.8	6.7978	23	00	00	00	9.2	621.8	6.0500	24	18	20	00	3.4	621.8	5.0648		
25	10	30	00	8.4	735.8	6.7978	23	00	45	00	7.2	621.8		24	18	37	00	3.8	621.8			
25	11	35	00	8.5	735.8	7.0645	23	01	00	00	8.0	621.8		24	19	20	00	5.2	621.1	7.6258		
25	12	29	00	8.0	735.8	7.3403	23	02	45	00	7.2	622.6	6.5400	24	20	00	00	5.4	621.1	5.8176		
25	13	30	00	7.9	735.8	7.0645	23	03	00	00	7.7	622.6		24	20	15	00	6.3	621.1			
25	14	25	00	7.1	735.8	7.6258	23	03	00	00	7.3	622.6		24	20	30	00	7.1	620.3			
25	15	30	00	8.9	736.6	7.6258	23	04	00	00	6.2	622.6		24	20	40	00	7.3	620.3	6.4397		
25	16	17	00	11.1	737.3	8.8679	23	04	45	00	6.2	622.6		24	20	50	00	7.5	620.3			
25	17	15	00	12.3	738.1	9.2050	23	05	00	00	6.5	622.6		24	21	00	00	8.8	620.3			
25	18	16	00	14.1	737.3	10.2850	23	06	00	00	6.5	622.6	5.4836	24	21	05	00	7.5	619.6			
25	19	24	00	15.1	735.8	10.2650	23	06	40	00	6.8	622.6	6.2908	24	21	15	00	7.9	619.6			
25	20	20	00	16.0	735.1	10.6688	23	07	00	00	7.0	622.6		24	22	00	00	8.7	619.6			
25	21	24	00	16.0	735.1	9.9131	23	08	00	00	7.3	621.8		24	22	15	00	8.9	619.6	7.0645		
25	22	20	00	16.9	734.3	10.6688	23	08	40	00	7.7	621.1	4.9647	24	22	20	00	8.8	619.6			
25	23	26	00	15.3	734.3	10.2850	23	09	00	00	8.0	621.1		24	22	23	00	9.4	619.6			
25	24	20	00	14.4	735.1	9.9131	23	09	00	00	8.0	621.1		24	22	32	00	7.3	619.6	7.2012		
							23	10	00	00	4.8	621.1		24	22	35	00	7.5	619.6			
							23	11	00	00	5.1	621.1		24	22	45	00	6.9	619.6	6.6937		
							23	12	00	00	4.4	621.1		24	23	15	00	5.4	619.6			
							23	12	45	00	4.6	621.1		24	23	35	00	5.0	619.6			
							23	19	30	00	4.9	622.6	5.3762	25	00	15	00	4.3	618.8	5.0648		
							23	20	00	00	5.1	622.6		25	00	20	00	4.0	619.6			
							23	21	00	00	5.5	621.8		25	00	25	00	3.8	619.6			
							23	22	00	00	4.2	621.8		25	00	29	00	3.9	618.8			
							23	23	00	00	2.2	621.8		25	01	15	00	2.8	618.8			
							24	00	00	00	1.4	621.8		25	01	20	00	2.7	618.8			
							24	01	45	00	1.0	621.8		25	02	15	00	2.4	618.8	4.7696		
							24	02	45	00	1.0	621.8		25	02	27	00	2.3	618.8			
							24	03	45	00	1.0	622.6		25	02	32	00	2.5	618.8			
							24	04	45	00	.8	622.6	4.3764	25	02	39	00	2.3	618.8			
							Santiago															
							-----															
							11.0	623.3														
							10.3	623.3														
							10.1	623.3														
							10.2	623.3														
							12.4	623.3														
							10.2	622.6														
							10.5	623.3														
							10.3	623.3														
							10.3	623.3														
							10.0	623.3														
							9.8	623.3														

Table A2. (continued)

Time		T	P	V	Time		T	P	V	Time		T	P	V
dy	hr mn sc	deg C	mmHg	mmHg	dy	hr mn sc	deg C	mmHg	mmHg	dy	hr mn sc	deg C	mmHg	mmHg
25 02	43 00	2.3	618.8		26 00	20 00	5.5	617.3	6.1694	24 04	45 00	8.5	719.3	8.3828
25 02	51 00	2.4	618.8	4.7315						24 05	00 00	8.5	719.3	8.2263
25 03	00 00	2.3	618.8							24 06	30 00	8.0	719.3	8.0724
25 03	15 00	1.7	618.8							24 06	37 00	8.0	719.3	7.8313
25 03	27 00	1.5	618.8							24 08	28 00	8.6	719.3	7.6840
25 04	15 00	1.1	618.8	4.5811						24 10	05 00	9.0	718.6	7.6258
25 04	25 00	1.0	618.8							24 11	01 00	8.0	718.6	7.3403
25 05	15 00	.5	618.8							24 12	00 00	8.0	718.6	7.3403
25 05	25 00	.8	618.8							24 13	10 00	8.0	718.6	7.1737
25 06	15 00	1.8	618.8	4.5811						24 14	00 00	7.5	717.8	7.0545
25 06	23 00	2.1	618.8							24 15	02 00	7.0	717.8	7.0645
25 06	26 00	2.0	618.8							24 16	03 00	8.0	718.6	7.4534
25 06	35 00	1.9	618.8							24 16	33 00	8.6	718.6	7.6548
25 06	37 00	1.9	618.8							24 17	31 00	11.5	718.6	8.4778
25 06	50 00	2.1	618.8							24 18	29 00	13.2	717.8	9.7318
25 07	15 00	2.3	618.8							24 19	34 00	15.0	717.1	9.8402
25 07	25 00	2.8	618.1	4.3764						24 20	30 00	16.5	716.3	9.7318
25 08	15 00	.6	618.1							24 21	30 00	15.5	716.3	9.6350
25 08	20 00	.2	618.1							24 22	30 00	16.0	715.5	9.5332
25 08	24 00	.1	618.1							24 23	30 00	14.5	715.6	9.0350
25 09	20 00	1.1	617.3							25 00	30 00	12.0	715.6	8.9724
25 10	15 00	1.9	617.3							25 01	30 00	9.5	715.6	8.0724
25 10	20 00	1.8	617.3	4.2774						25 02	36 00	8.5	715.6	7.9210
25 10	23 00	1.8	617.3							25 03	30 00	8.0	715.6	7.6258
25 10	35 00	1.6	617.3							25 04	35 00	8.5	715.6	7.3610
25 10	38 00	1.7	617.3							25 05	30 00	9.5	715.6	8.2885
25 10	50 00	1.8	617.3							25 06	35 00	8.2	715.6	7.0645
25 11	15 00	2.1	617.3							25 07	30 00	7.5	714.8	7.2012
25 11	20 00	2.1	617.3							25 08	32 00	8.0	714.1	6.6678
25 12	15 00	.8	616.6	3.8100						25 09	30 00	9.0	713.3	6.3894
25 12	20 00	.3	616.6	4.3764						25 10	31 00	10.0	713.3	6.1935
25 12	23 00	.4	616.6							25 11	30 00	7.5	713.3	6.6678
25 13	15 00	.6	617.3							25 12	31 00	7.5	713.3	6.7978
25 13	20 00	.5	617.3							25 13	35 00	7.0	714.1	6.4354
25 14	20 00	-.3	617.3	4.3764						25 14	29 00	7.0	714.1	6.0028
25 15	20 00	1.9	618.1							25 15	32 00	9.3	714.8	7.5969
25 16	20 00	4.7	618.1							25 16	32 00	11.0	715.6	8.5058
25 17	20 00	5.5	618.8	6.0500						25 17	33 00	17.0	715.6	9.2050
25 18	20 00	6.2	618.8							25 18	33 00	14.4	714.8	8.5418
25 19	20 00	7.7	618.1	6.2908						25 19	30 00	15.5	714.1	8.2263
25 20	20 00	8.3	617.3							25 20	35 00	18.0	713.5	8.3828
25 21	20 00	9.4	617.3							25 21	45 00	18.5	712.6	7.9210
25 22	20 00	8.8	617.3							25 22	33 00	20.0	712.6	7.6840
25 23	20 00	6.8	617.3							26 00	35 00	16.5	713.3	7.4818

San Juan

Time			n			Time			n			Time			n		
dy	hr	mn	sc	n	Time	dy	hr	mn	sc	n	Time	dy	hr	mn	sc	n	
Mt. Wilson																	
22	12	08	01	1.00024386		24	13	54	08	1.00025008		26	01	08	09	1.00024528	
22	13	08	54	1.00024473		24	14	28	00	1.00025008		27	21	13	23	1.00024919	
22	14	53	40	1.00024414		24	14	54	16	1.00025008		Castro					
22	15	59	57	1.00024414		24	15	53	54	1.00025040		-----					
22	16	31	16	1.00024327		24	16	41	49	1.00024858		22	09	10	00	1.00026392	
22	16	57	27	1.00024414		24	17	53	26	1.00024796		22	10	00	00	1.00026383	
22	17	59	44	1.00024272		24	19	06	29	1.00024679		22	11	00	00	1.00026448	
22	19	14	19	1.00024071		24	19	55	54	1.00024679		22	12	00	00	1.00026465	
22	21	57	23	1.00024183		24	20	49	38	1.00024558		22	13	00	00	1.00026457	
22	22	22	25	1.00024129		24	21	54	25	1.00024411		22	14	00	00	1.00026493	
22	23	02	15	1.00024183		24	22	52	51	1.00024411		22	15	00	00	1.00026502	
22	23	00	00	1.00024269		24	23	55	10	1.00024588		22	16	00	00	1.00026447	
22	23	00	28	1.00024355		25	00	34	16	1.00024499		22	17	00	00	1.00026459	
23	01	03	14	1.00024355		25	02	00	11	1.00024766		22	18	00	00	1.00026376	
23	02	01	37	1.00024442		25	03	03	19	1.00024857		22	19	00	00	1.00026294	
23	02	24	05	1.00024442		25	03	55	51	1.00024948		22	20	00	00	1.00025994	
23	02	48	26	1.00024529		25	04	32	31	1.00024857		22	21	00	00	1.00026228	
23	03	04	13	1.00024618		25	05	54	44	1.00024857		22	22	00	00	1.00026328	
23	03	05	57	1.00024649		25	06	28	44	1.00024857		22	23	00	00	1.00026237	
23	03	06	37	1.00024618		25	07	01	06	1.00024857		23	01	00	00	1.00026411	
23	03	06	55	1.00024529		25	07	56	42	1.00024919		23	01	00	00	1.00026522	
23	03	07	52	1.00024590		25	08	36	19	1.00024887		23	02	00	00	1.00026632	
23	03	08	20	1.00024590		25	09	48	16	1.00024796		23	03	00	00	1.00026764	
23	03	08	54	1.00024590		25	10	28	22	1.00024706		23	04	00	00	1.00026615	
23	03	09	56	1.00024679		25	11	00									

**San Pedro**

Table A3. (continued)

Time				Time				Time				Time							
dy	hr	mn	sc	n	dy	hr	mn	sc	n	dy	hr	mn	sc	n	dy	hr	mn	sc	n
22	12	20	00	1.00027576	22	19	33	00	1.00027924	24	19	15	00	1.00028302	22	14	49	00	1.00024780
22	12	20	00	1.00027684	22	20	32	00	1.00027792	24	20	23	00	1.00028069	22	15	00	00	1.00024166
22	12	20	00	1.00027397	22	21	40	00	1.00028002	24	21	40	00	1.00028033	22	16	00	00	1.00024214
22	12	18	20	1.00027095	22	22	30	00	1.00028137	24	22	30	00	1.00028391	22	16	45	00	1.00024214
22	12	22	00	1.00027076	22	23	37	00	1.00028305	24	23	30	00	1.00028101	22	17	00	00	1.00024214
23	10	20	00	1.00027733	23	00	38	00	1.00028384	25	00	27	00	1.00028327	22	18	00	00	1.00024146
23	16	20	00	1.00027760	23	01	45	00	1.00028534	25	01	15	00	1.00028536	22	18	45	00	1.00023960
23	18	20	00	1.00027731	23	02	32	00	1.00028464	25	02	27	00	1.00028698	22	19	00	00	1.00023971
23	20	30	00	1.00027509	23	03	35	00	1.00028534	25	03	30	00	1.00028600	22	20	00	00	1.00023971
24	00	20	00	1.00027811	23	04	40	00	1.00028524	25	04	23	00	1.00028851	22	21	00	00	1.00023903
24	04	20	00	1.00028055	23	05	30	00	1.00028534	25	05	40	00	1.00028728	22	22	00	00	1.00024020
24	06	10	00	1.00028025	23	06	32	00	1.00028543	25	06	25	00	1.00028717	22	22	45	00	1.00024150
24	06	45	00	1.00028035	23	07	30	00	1.00028533	25	07	28	00	1.00028758	22	23	00	00	1.00024071
24	08	10	00	1.00028038	23	08	30	00	1.00028228	25	08	23	00	1.00028711	22	23	45	00	1.00024251
24	10	10	00	1.00028107	23	09	29	00	1.00028393	25	09	29	00	1.00028720	23	00	00	00	1.00024193
24	10	45	00	1.00028117	23	10	35	00	1.00028423	25	10	30	00	1.00028710	23	00	45	00	1.00024356
24	12	10	00	1.00028077	23	11	30	00	1.00028423	25	11	35	00	1.00028700	23	01	00	00	1.00024396
24	14	10	00	1.00028097	23	12	40	00	1.00028413	25	12	29	00	1.00028751	23	02	00	00	1.00024397
24	14	45	00	1.00028127	23	17	40	00	1.00028469	25	13	30	00	1.00028761	23	02	45	00	1.00024354
25	02	10	00	1.00027927	23	18	34	00	1.00028429	25	14	25	00	1.00028843	23	03	00	00	1.00024398
25	02	52	00	1.00027927	23	19	30	00	1.00028429	25	15	30	00	1.00028690	23	04	00	00	1.00024434
25	04	10	00	1.00027927	23	20	40	00	1.00028128	25	16	17	00	1.00028495	23	04	45	00	1.00024494
25	06	10	00	1.00027998	23	21	33	00	1.00028157	25	17	15	00	1.00028406	23	05	00	00	1.00024453
25	06	46	00	1.00028019	23	22	30	00	1.00028205	25	18	16	00	1.00028198	23	06	00	00	1.00024450
25	08	10	00	1.00028099	23	23	26	00	1.00028137	25	19	24	00	1.00028043	23	06	45	00	1.00024432
25	10	10	00	1.00027999	24	00	30	00	1.00028452	25	20	20	00	1.00027929	23	07	00	00	1.00024414
25	10	45	00	1.00027999	24	01	40	00	1.00028679	25	21	24	00	1.00027929	23	08	00	00	1.00024557
25	12	10	00	1.00028099	24	02	40	00	1.00028710	25	22	20	00	1.00027912	23	08	45	00	1.00024595
25	14	10	00	1.00028067	24	03	50	00	1.00028699	25	22	26	00	1.00027955	23	08	45	00	1.00024595
25	14	22	00	1.00028119	24	04	35	00	1.00028710	25	23	20	00	1.00027955	23	09	00	00	1.00024595
25	16	10	00	1.00027770	24	05	40	00	1.00028750	26	00	20	00	1.00028084	23	10	00	00	1.00024549
										Santiago									
										22	09	00	00	1.00024098	23	12	00	00	1.00024522
										22	10	00	00	1.00024157	23	12	45	00	1.00024534
										22	10	40	00	1.00024174	23	13	00	00	1.00024566
										22	10	46	00	1.00024166	23	19	30	00	1.00024599
										22	11	00	00	1.00023980	23	20	00	00	1.00024581
										22	12	00	00	1.00024139	23	21	00	00	1.00024514
										22	12	42	00	1.00024140	23	22	00	00	1.00024529
										22	12	46	00	1.00024157	23	23	00	00	1.00024606
										22	13	00	00	1.00024157	24	00	00	00	1.00024600
										22	13	00	00	1.00024183	24	01	45	00	1.00024917
										22	14	00	00	1.00024260	24	02	45	00	1.00024917
										22	18	17	00	1.00028492	24	03	45	00	1.00024949
										22	18	31	00	1.00028020	24	04	45	00	1.00024967
										Niguel									
										22	09	28	00	1.00028344	23	12	00	00	1.00024522
										22	10	32	00	1.00028413	23	12	45	00	1.00024534
										22	11	30	00	1.00028463	23	19	30	00	1.00024599
										22	12	30	00	1.00028523	23	20	00	00	1.00024581
										22	13	31	00	1.00028513	23	21	00	00	1.00024514
										22	14	30	00	1.00028286	23	22	00	00	1.00024529
										22	15	41	00	1.00028362	23	23	00	00	1.00024606
										22	16	36	00	1.00028224	24	00	00	00	1.00024600
										22	17	30	00	1.00028107	24	01	45	00	1.00024917
										22	18	31	00	1.00028020	24	02	45	00	1.00024917

Table A3. (continued)

Time		n	Time		n	Time		n	Time		n
dy	hr	mn	sc	dy	hr	mn	sc	dy	hr	mn	sc
24	05	45	00	1.00024949	1.00024679	1.00024300	25	02	43	00	1.00024337
24	06	25	00	1.00024940	1.00024670	1.00024300	25	02	51	00	1.00024337
24	06	40	00	1.00024931	1.00024679	1.00024300	25	03	00	00	1.00024337
24	07	40	00	1.00024944	1.00024733	1.00024300	25	03	15	00	1.00024337
24	08	20	00	1.00024935	1.00024751	1.00024300	25	03	27	00	1.00024337
24	09	20	00	1.00024962	1.00024787	1.00024300	25	04	15	00	1.00024337
24	10	20	00	1.00024907	1.00024796	1.00024300	25	04	25	00	1.00024337
24	11	20	00	1.00024943	1.00024842	1.00024300	25	05	15	00	1.00024337
24	12	20	00	1.00024907	1.00024815	1.00024300	25	05	25	00	1.00024337
24	13	20	00	1.00024907	1.00024724	1.00024300	25	06	15	00	1.00024337
24	14	20	00	1.00024816	1.00024697	1.00024300	25	06	23	00	1.00024337
24	15	20	00	1.00024690	1.00024706	1.00024300	25	06	26	00	1.00024337
24	16	20	00	1.00024780	1.00024715	1.00024300	25	06	35	00	1.00024337
24	17	20	00	1.00024692	1.00024715	1.00024300	25	06	37	00	1.00024337
24	18	20	00	1.00024700	1.00024697	1.00024300	25	06	50	00	1.00024337
24	18	37	00	1.00024665	1.00024679	1.00024300	25	07	15	00	1.00024337
24	19	20	00	1.00024513	1.00024607	1.00024300	25	07	25	00	1.00024337
24	20	00	00	1.00024495	1.00024607	1.00024300	25	08	15	00	1.00024337
24	20	15	00	1.00024417	1.00024605	1.00024300	25	08	20	00	1.00024337
24	20	30	00	1.00024316	1.00024650	1.00024300	25	08	24	00	1.00024337
24	20	40	00	1.00024298	1.00024727	1.00024300	25	09	20	00	1.00024337
24	20	50	00	1.00024281	1.00024655	1.00024300	25	10	15	00	1.00024337
24	21	00	00	1.00024169	1.00024664	1.00024300	25	10	20	00	1.00024337
24	21	05	00	1.00024253	1.00024664	1.00024300	25	10	23	00	1.00024337
24	21	15	00	1.00024219	1.00024692	1.00024300	25	10	35	00	1.00024337
24	22	00	00	1.00024150	1.00024673	1.00024300	25	10	38	00	1.00024337
24	22	15	00	1.00024133	1.00024664	1.00024300	25	10	50	00	1.00024337
24	22	20	00	1.00024142	1.00024638	1.00024300	25	11	15	00	1.00024337
24	22	23	00	1.00024090	1.00024638	1.00024300	25	11	20	00	1.00024337
24	22	32	00	1.00024228	1.00024726	1.00024300	25	12	15	00	1.00024337
24	22	35	00	1.00024253	1.00024772	1.00024300	25	12	20	00	1.00024337
24	22	45	00	1.00024305	1.00024762	1.00024300	25	12	23	00	1.00024337
24	23	15	00	1.00024436	1.00024772	1.00024300	25	13	15	00	1.00024337
24	23	35	00	1.00024471	1.00024782	1.00024300	25	13	20	00	1.00024337
25	00	15	00	1.00024502	1.00024854	1.00024300	25	14	20	00	1.00024337
25	00	20	00	1.00024560	1.00024687	1.00024300	25	15	20	00	1.00024337
25	00	25	00	1.00024578	1.00024439	1.00024300	25	16	20	00	1.00024337
25	00	29	00	1.00024537	1.00024396	1.00024300	25	17	20	00	1.00024337
25	01	15	00	1.00024635	1.00024335	1.00024300	25	18	20	00	1.00024337
25	01	20	00	1.00024644	1.00024178	1.00024300	25	19	20	00	1.00024337
25	02	15	00	1.00024670	1.00024095	1.00024300	25	20	20	00	1.00024337
25	02	27	00	1.00024679	1.00024001	1.00024300	25	21	20	00	1.00024337
25	02	32	00	1.00024662	1.00024052	1.00024300	25	22	20	00	1.00024337
25	02	39	00	1.00024679	1.00024224	1.00024300	25	23	20	00	1.00024337

### Table A4. Time-of-Flight Ratios

Castro			San Pedro			Niquel			Santiago			San Juan		
dy	hr	mn sc	dy	hr	mn sc	dy	hr	mn sc	dy	hr	mn sc	dy	hr	mn sc
1.104+			0.948+			1.364+			1.210+					
22 10 04 29	914198		22 10 18 55	870439		24 04 20 32	486752		22 10 41 41	567622		22 10 51 17	648866	0.735+
22 12 02 41	915170		22 12 21 52	871345		24 06 19 58	487793		22 12 43 11	567332		22 12 51 51	645234	
22 14 05 43	914546		22 14 15 03	870661		24 06 46 28	485966		22 14 34 32	566961		22 14 48 04	647619	
23 00 09 40	913745		22 18 23 39	871637		24 08 18 26	486255		22 22 49 55	566262		22 18 55 03	649189	
23 02 09 25	913337		24 04 20 40	870474		24 10 15 13	486647		23 00 41 06	566832		22 22 54 50	548337	
23 06 05 03	913747		24 06 16 19	871245		24 10 40 16	486957		23 04 33 04	566222		23 05 47 01	641875	
24 06 08 46	914336		24 06 51 04	870690		24 14 15 10	486632		23 06 39 34	565878		23 08 47 14	647233	
24 06 58 17	914280		24 08 09 40	871172		24 14 40 18	486914		23 08 39 44	565151		24 04 43 14	648393	
24 08 00 19	913737		24 10 10 03	870674		24 16 18 20	486108		24 04 40 08	566802		24 06 29 45	647995	
24 09 59 54	914011		24 10 45 34	870708		25 02 21 58	486409		24 06 25 43	566603		24 06 36 37	649039	
24 10 55 03	914255		24 12 10 04	870911		25 02 48 49	485806		24 06 41 31	565969		24 08 25 49	646657	
24 12 00 58	914457		24 14 11 30	870981		25 10 16 05	486836		24 08 22 17	565526		24 12 25 00	648343	
24 14 01 34	914142		24 14 45 60	870330		25 10 41 13	486512		24 10 21 13	566338		24 14 25 00	640772	
24 14 55 05	915268		25 02 13 18	870356					24 10 35 15	566618		24 14 38 04	649358	
24 16 01 03	915165		25 02 53 27	871173					24 12 20 03	566316		24 16 22 54	649338	
24 18 00 06	914538		25 04 09 53	869977					24 14 20 03	564893		25 02 33 03	648673	
24 18 57 39	914327		25 06 10 05	870052					24 14 35 11	564750		25 02 36 56	646634	
24 22 44 42	913891		25 06 45 43	870209					24 16 29 53	565410		25 04 25 60	650036	
25 02 05 25	914683		25 08 11 36	871274					24 20 18 04	565626		25 05 32 42	648312	
25 02 53 04	914826		25 10 10 16	871390					25 00 25 13	565089		25 06 23 56	647559	
25 04 00 52	913710		25 10 45 23	870584					25 02 29 30	565803		25 10 25 11	647534	
25 06 01 17	914789		25 12 10 30	871738					25 02 39 51	565623		25 10 30 03	647375	
25 06 55 21	915255		25 14 15 11	871417					25 04 22 37	566544		25 12 25 03	648233	
25 08 04 49	914402		25 14 40 42	871316					25 06 22 51	566702		25 18 31 55	647344	
25 10 01 53	914614								25 06 35 48	566195		25 20 25 59	648656	
25 10 55 08	915208								25 08 20 32	566493		25 22 28 50	648712	
25 12 04 05	914277								25 10 20 08	566462		25 60 32 50	648333	
25 14 51 54	915409								25 10 35 11	566062				
25 16 00 28	914909								25 12 21 02	565658				
25 18 00 19	914473								25 14 21 41	566251				
25 18 59 43	914613								25 14 35 02	566204				
25 20 07 42	914243								25 16 20 06	567037				
25 20 37 24	913952								25 18 39 42	566858				
25 20 49 58	914766								25 20 20 59	566413				
25 22 51 57	912706								25 22 37 54	565985				
26 00 05 27	912961								26 00 26 53	565337				



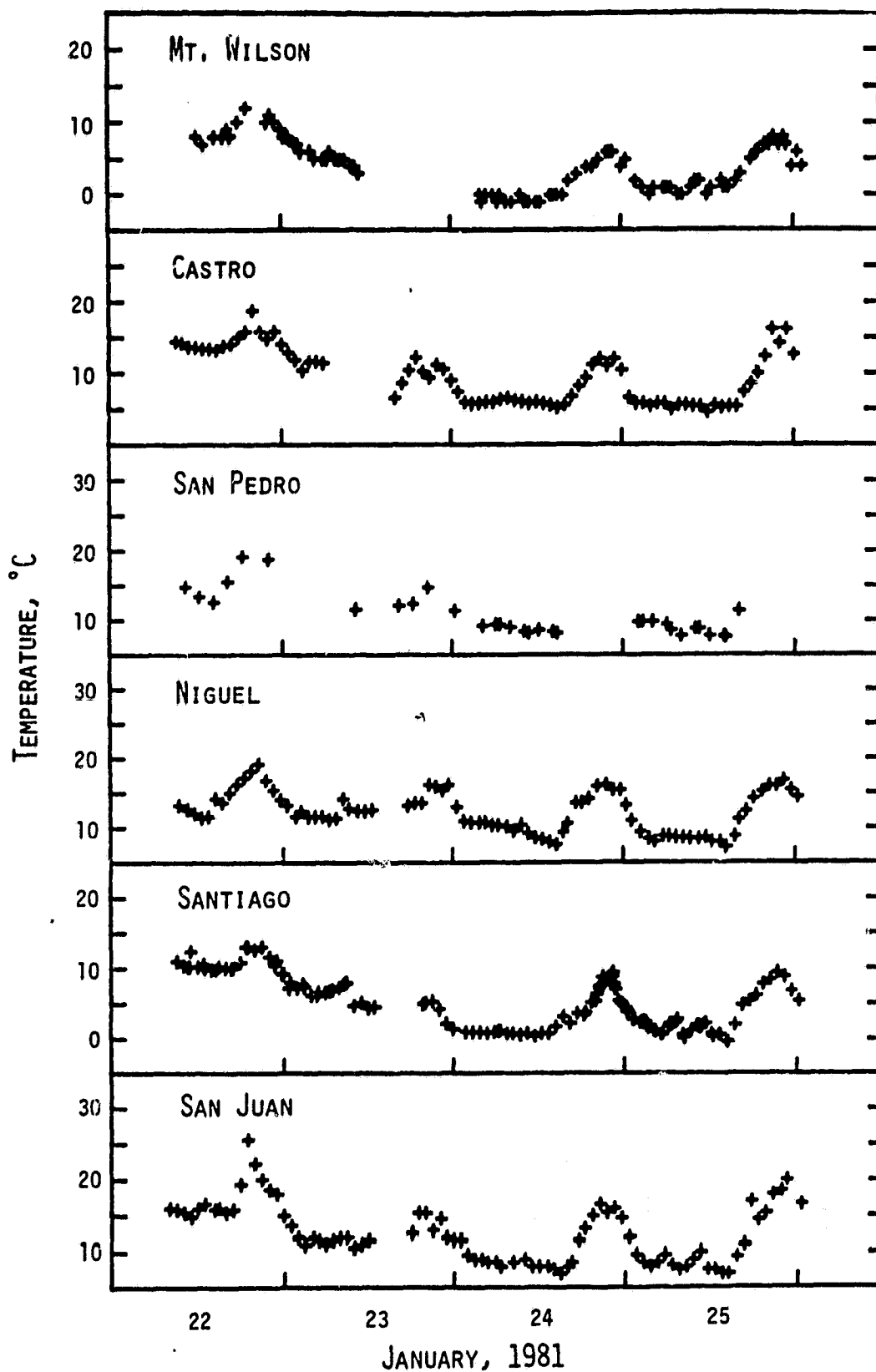


Fig. A1. Observed atmospheric temperature.

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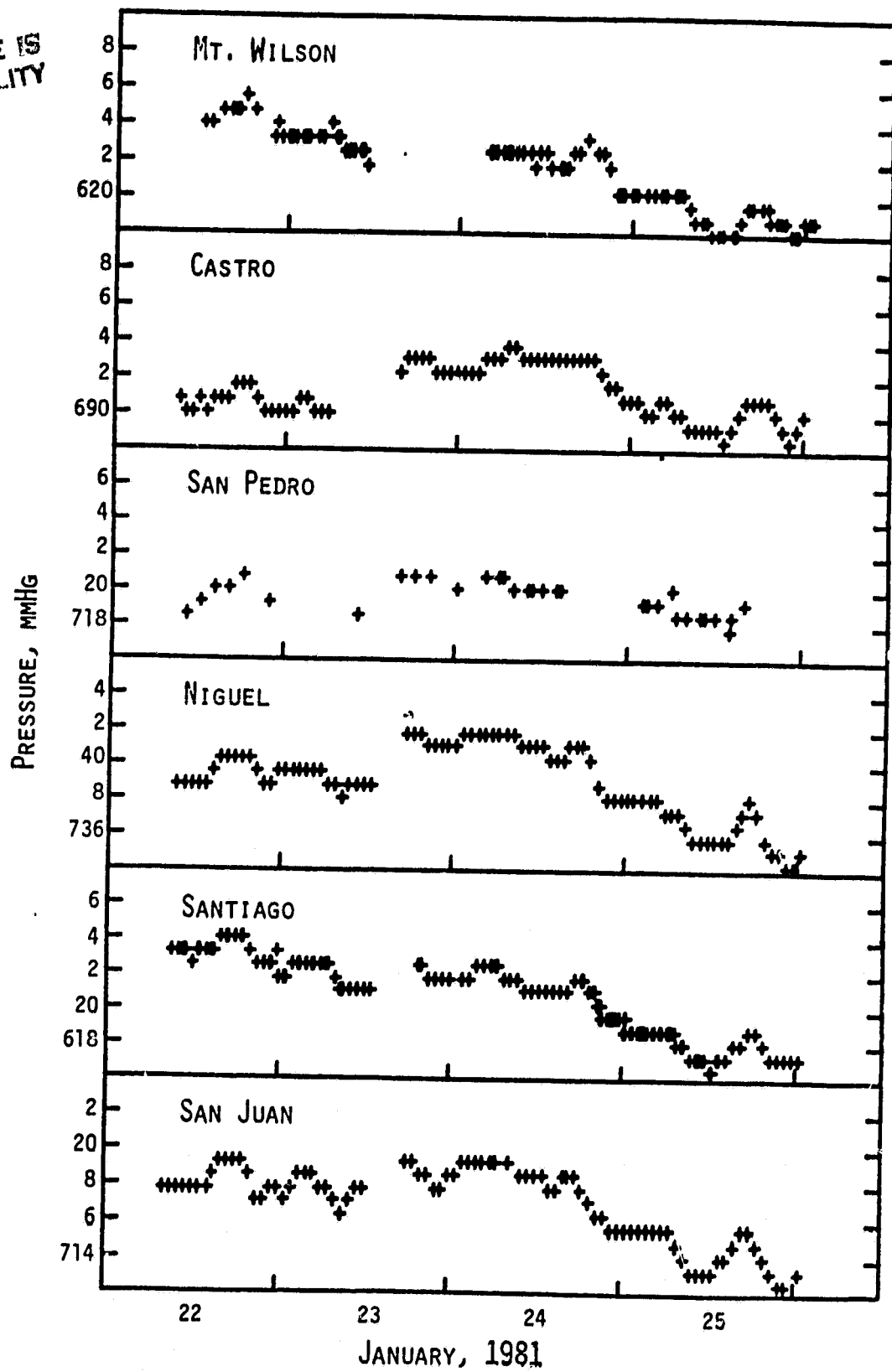


Fig. A2. Observed atmospheric pressure.

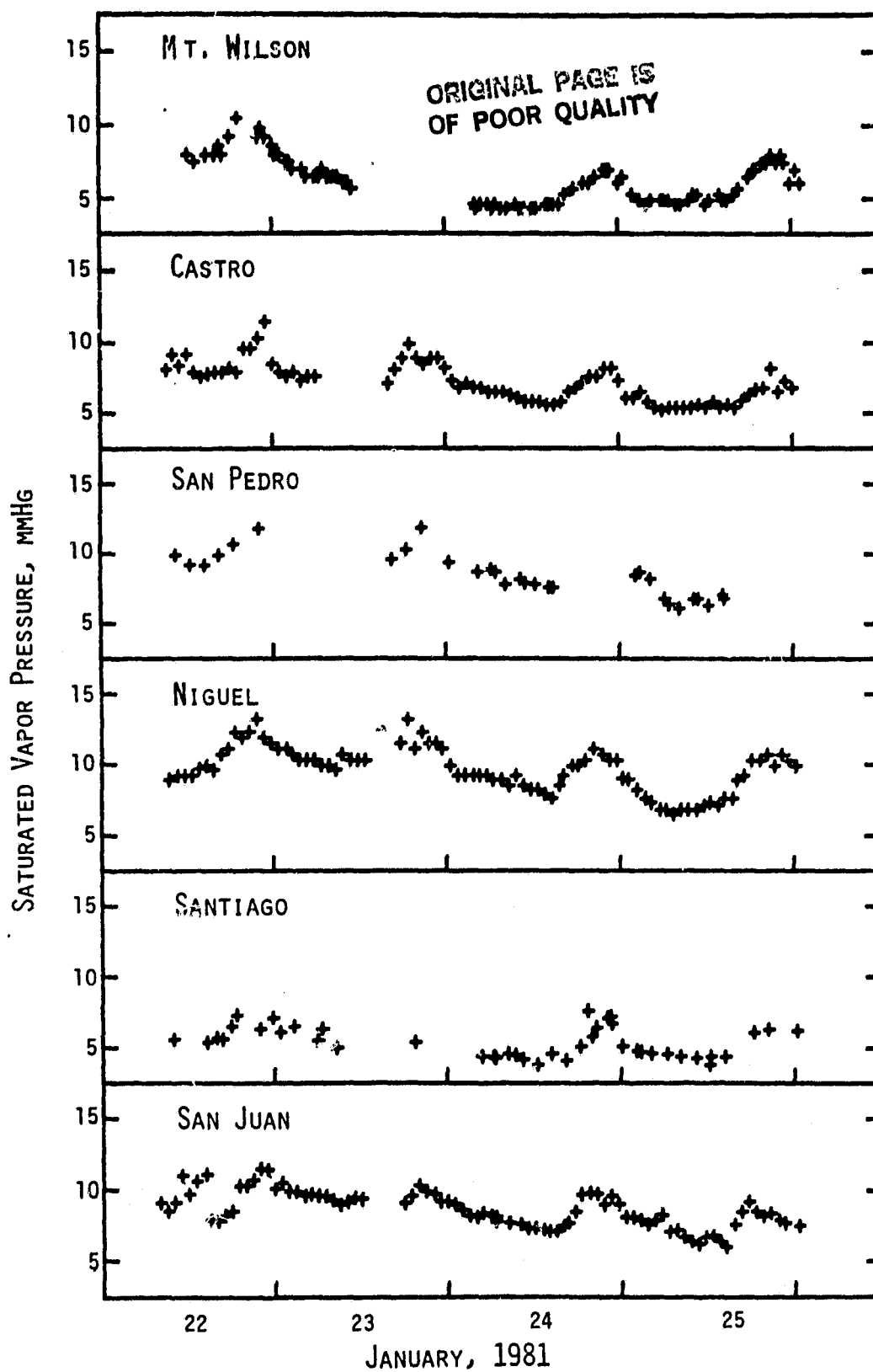


Fig. A3. Saturated vapor pressure.